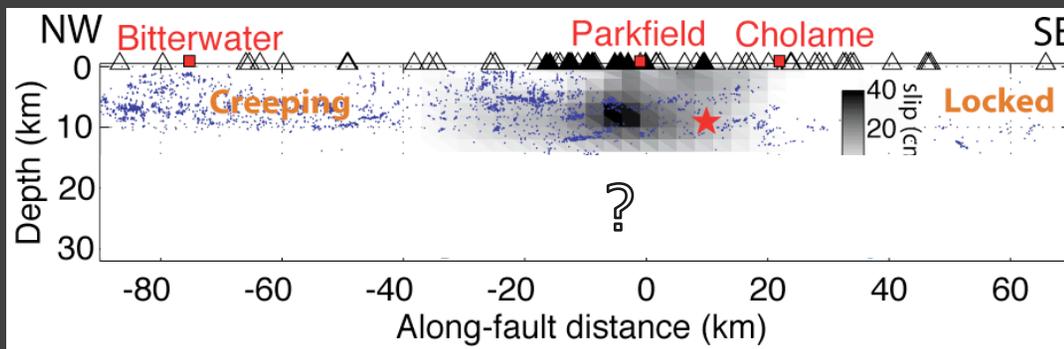


# Variations in Tremor Activity and Implications for Lower Crustal Deformation Along the Central San Andreas Fault



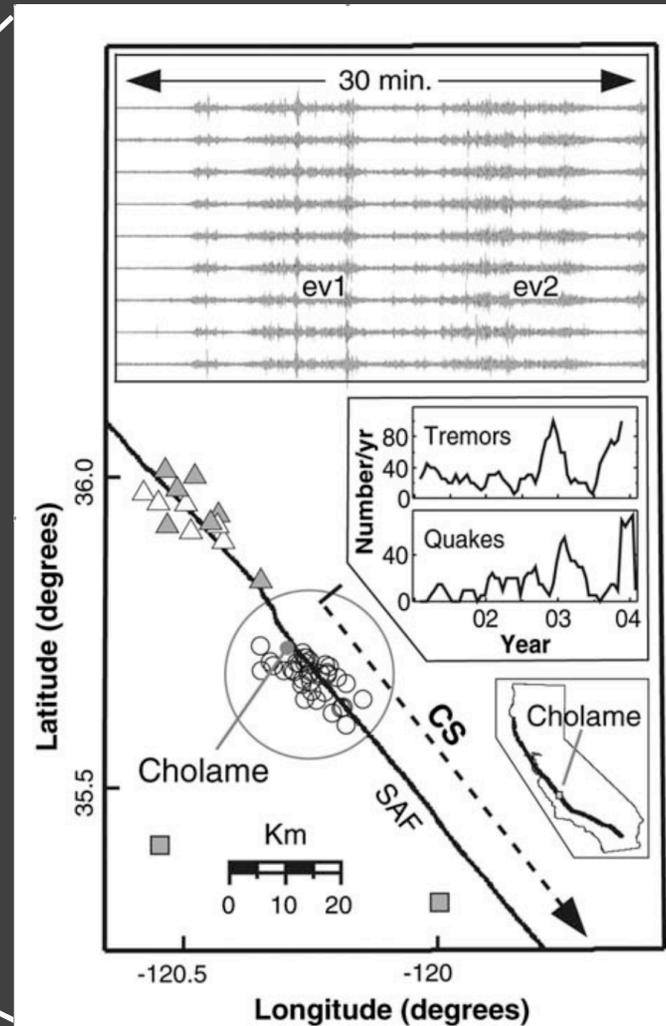
*Shelly and Hardebeck, GRL, 2010*

David R. Shelly  
USGS, Menlo Park



Collaborators:  
Jeanne Hardebeck  
Zhigang Peng  
Gary Fuis  
Jessica Murray-Moraleda  
Bill Ellsworth  
Roland Bürgmann  
Bob Nadeau  
Greg Beroza

# Tremor Under the San Andreas



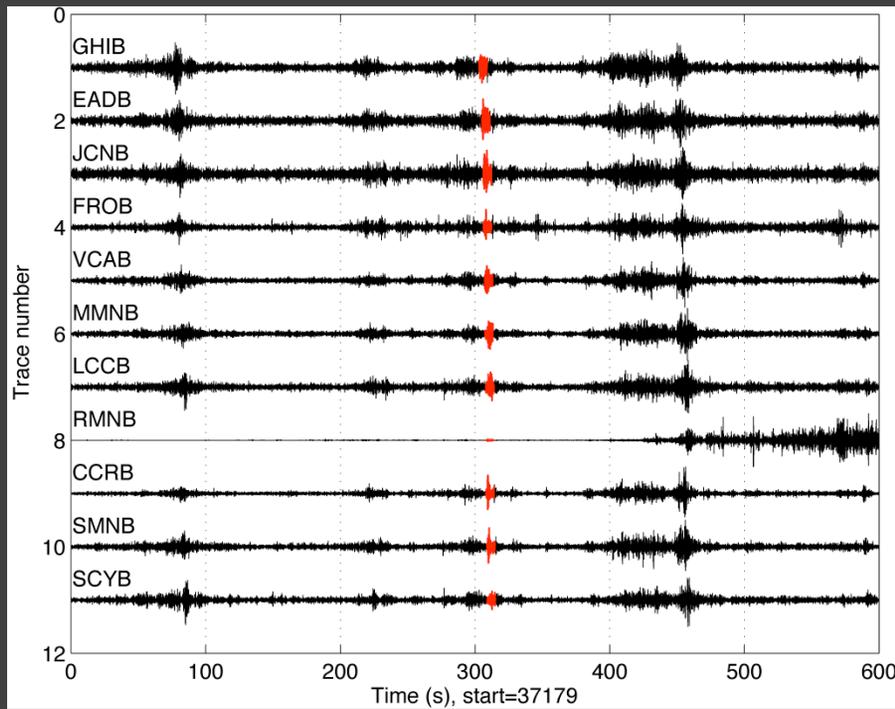
Nadeau and Dolenc, *Science*, 2005

# 1. Tremor Locations and Migration

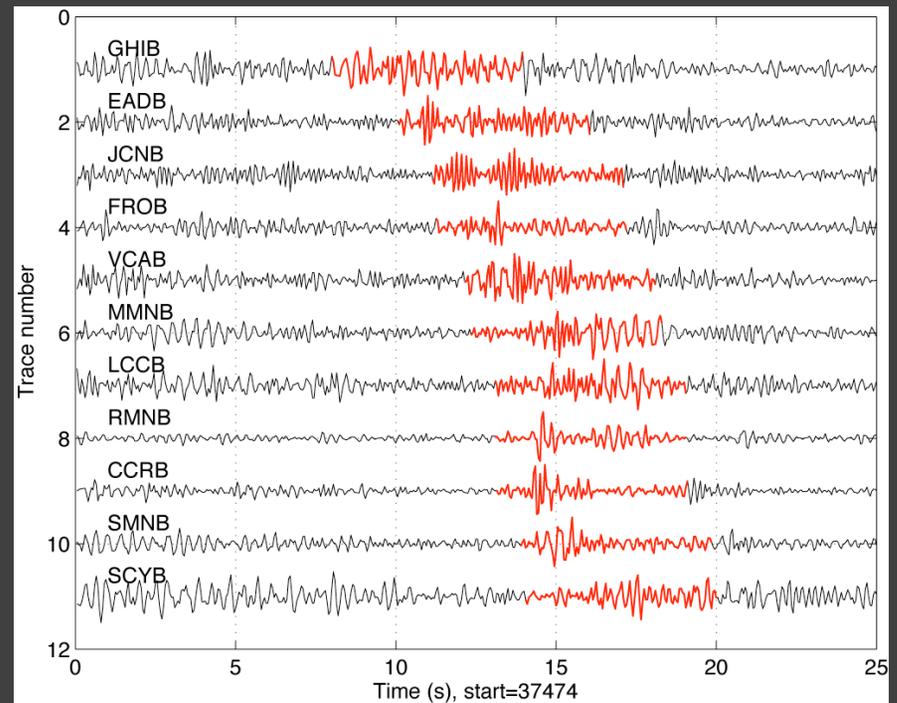
2 simple techniques:

- 1) Cross-correlation (multiplication)
- 2) Stacking (addition)

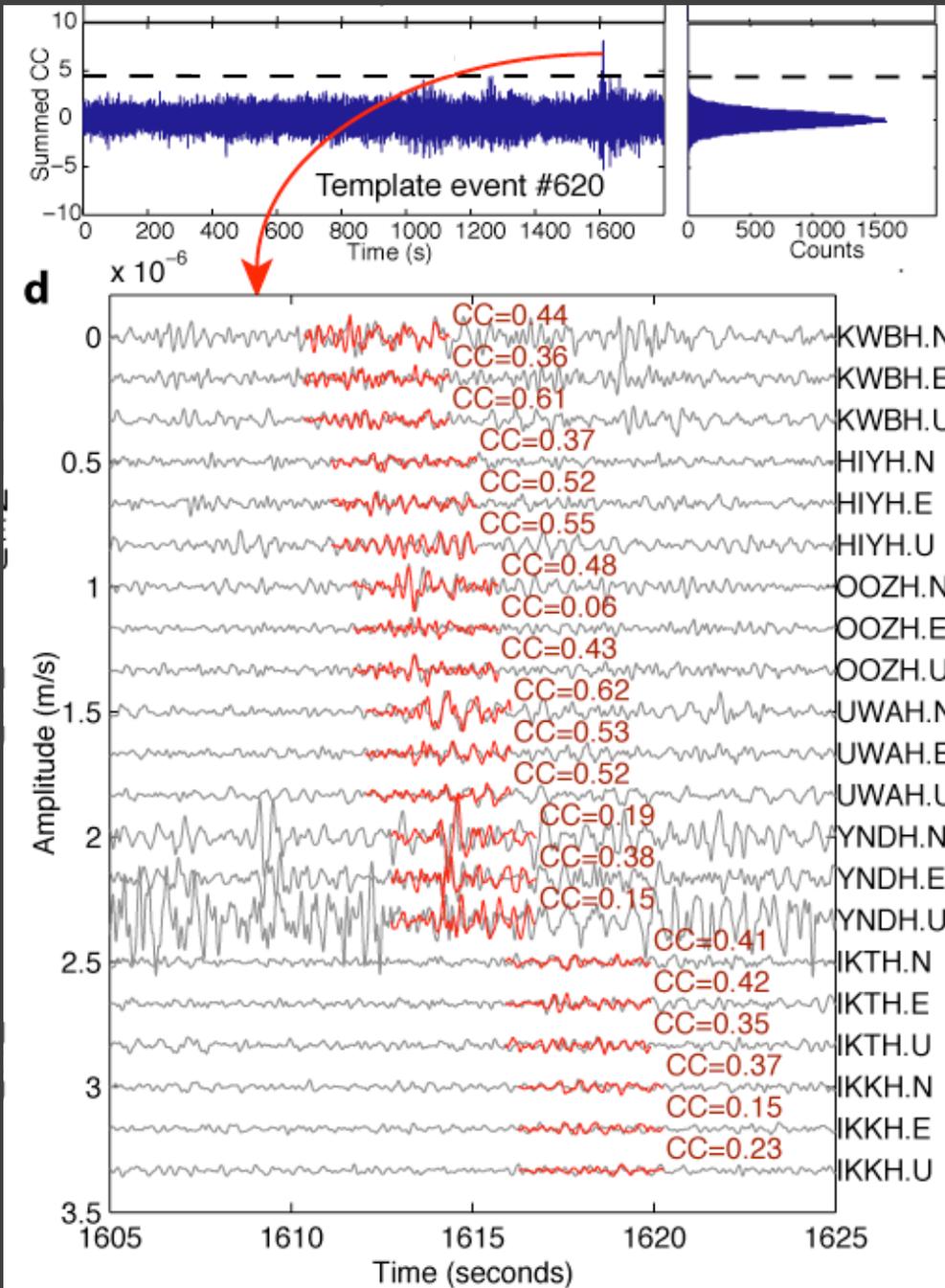
# Template Waveforms



10 minutes



25 seconds

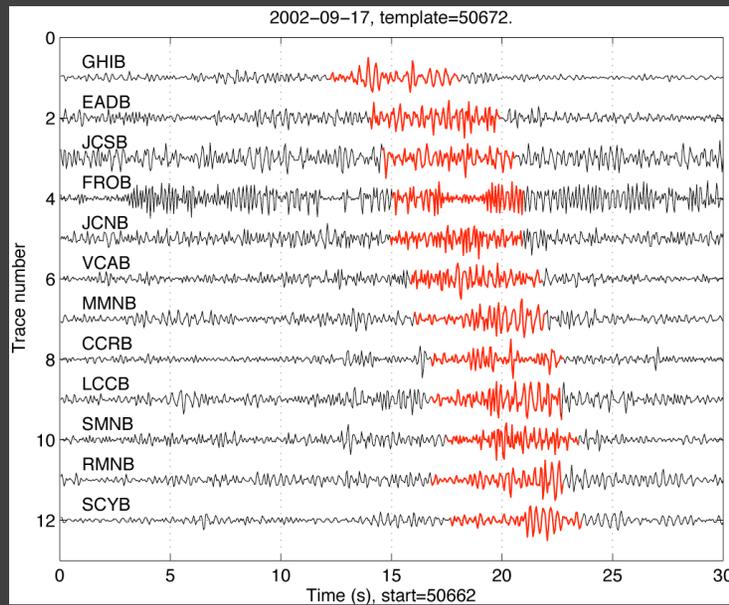


Scan template through continuous data, sum correlations

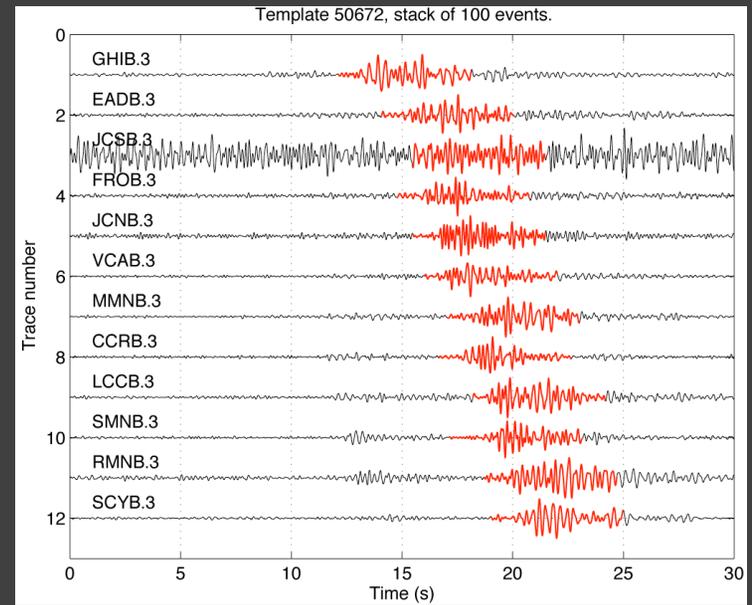
Shelly *et al.*, Nature, 2007

# Stacked LFE Templates

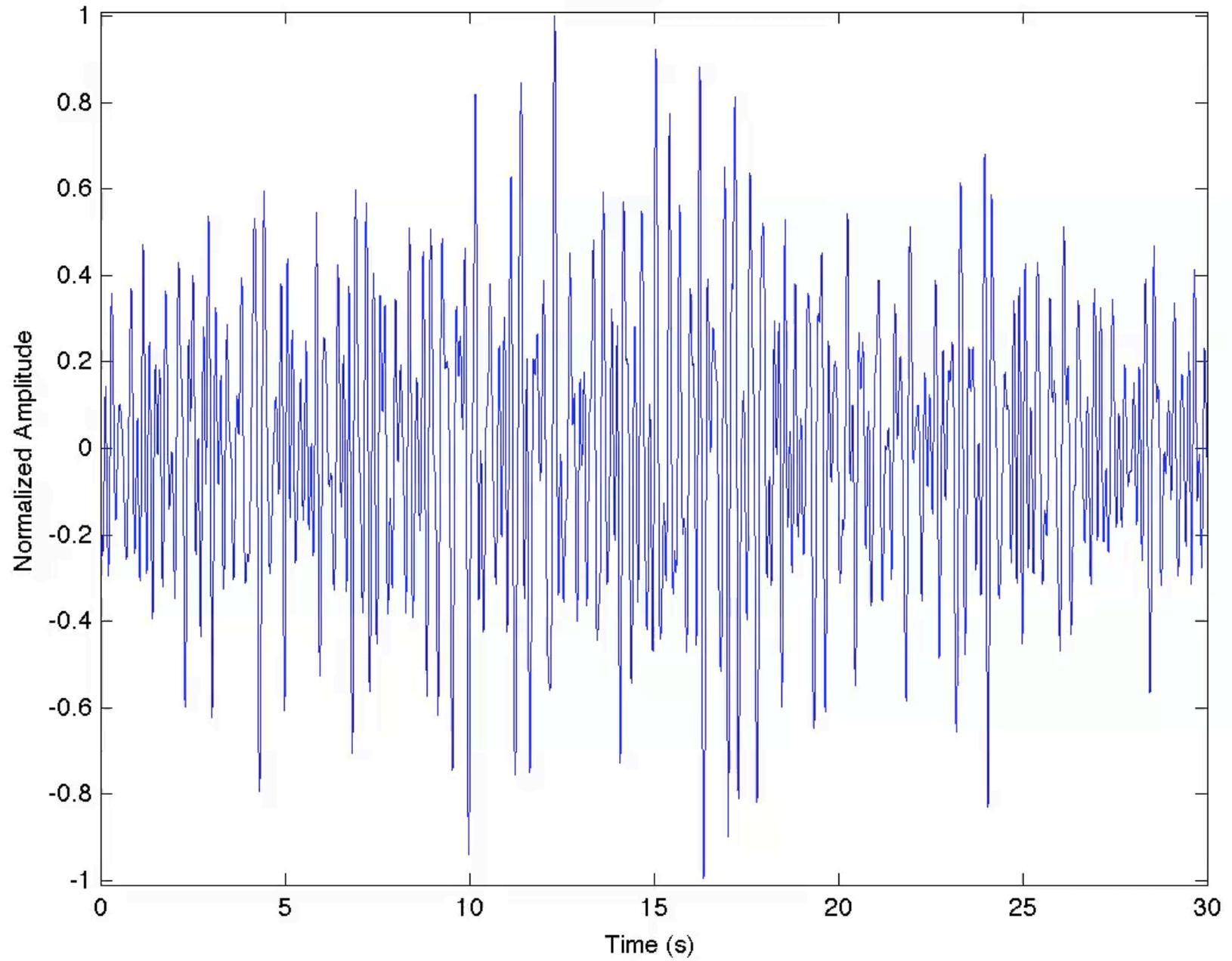
## Single LFE template



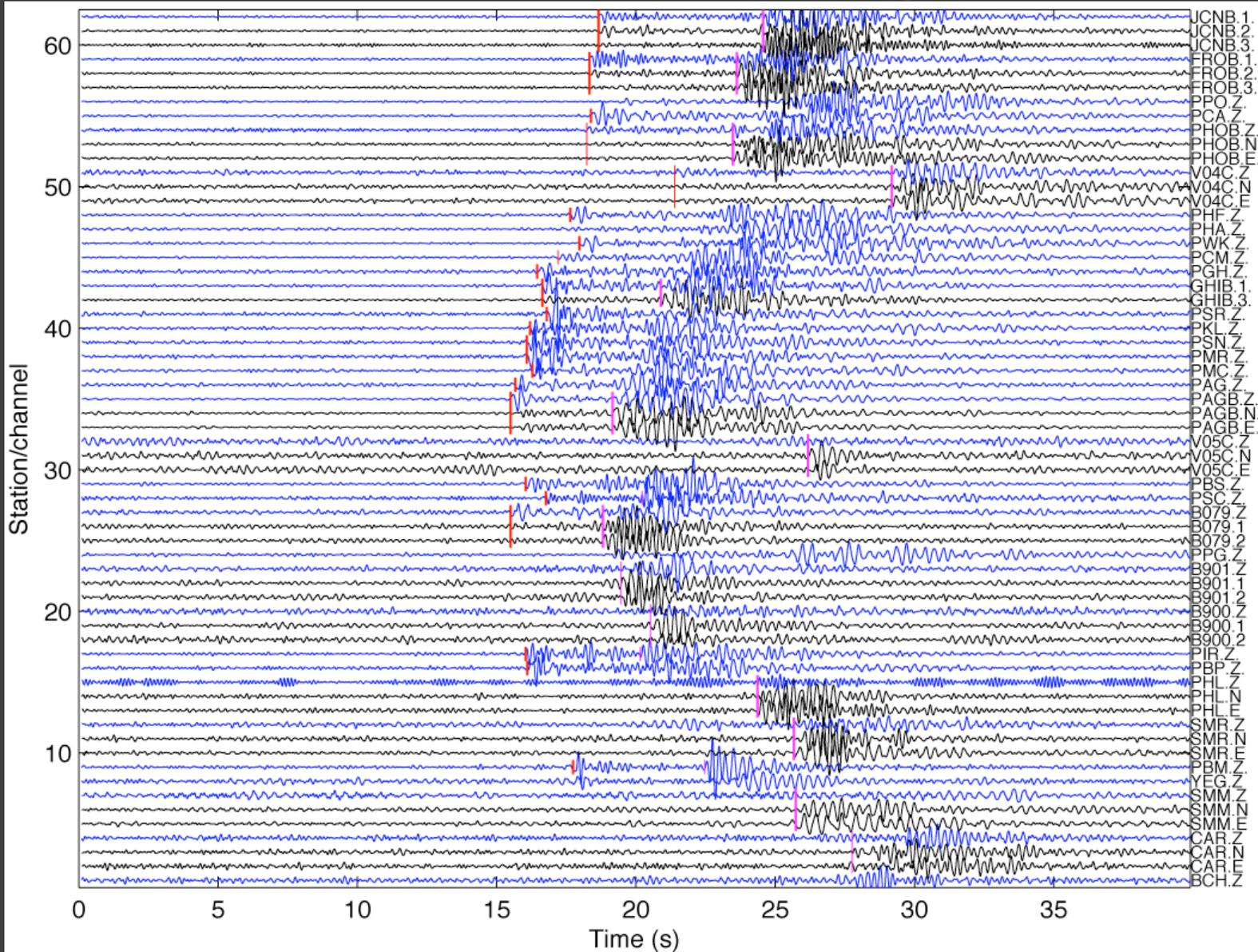
## 100 LFE stack template



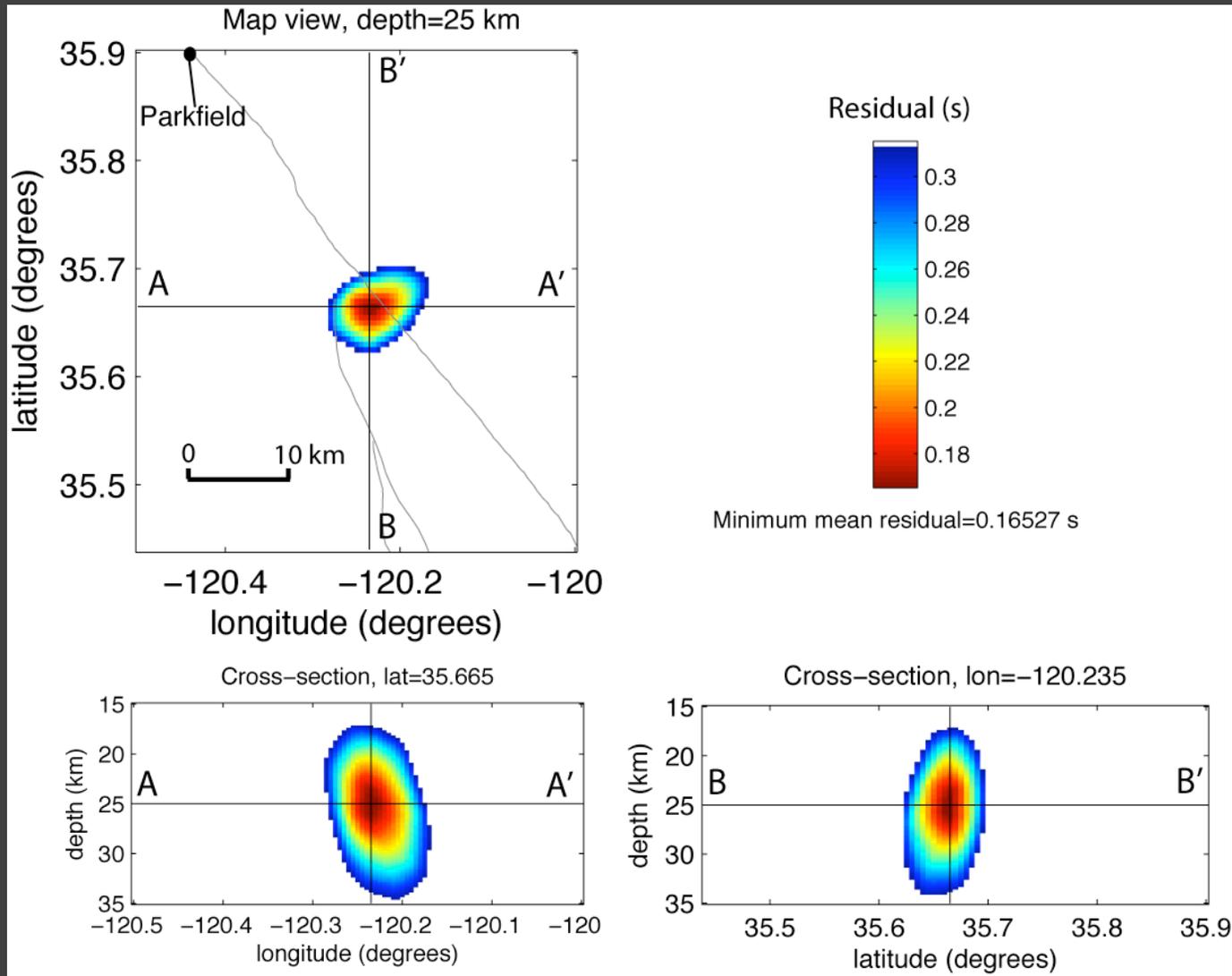
1 stacked events



# Stacked Waveforms/Picks



# Grid Search Location (3D)

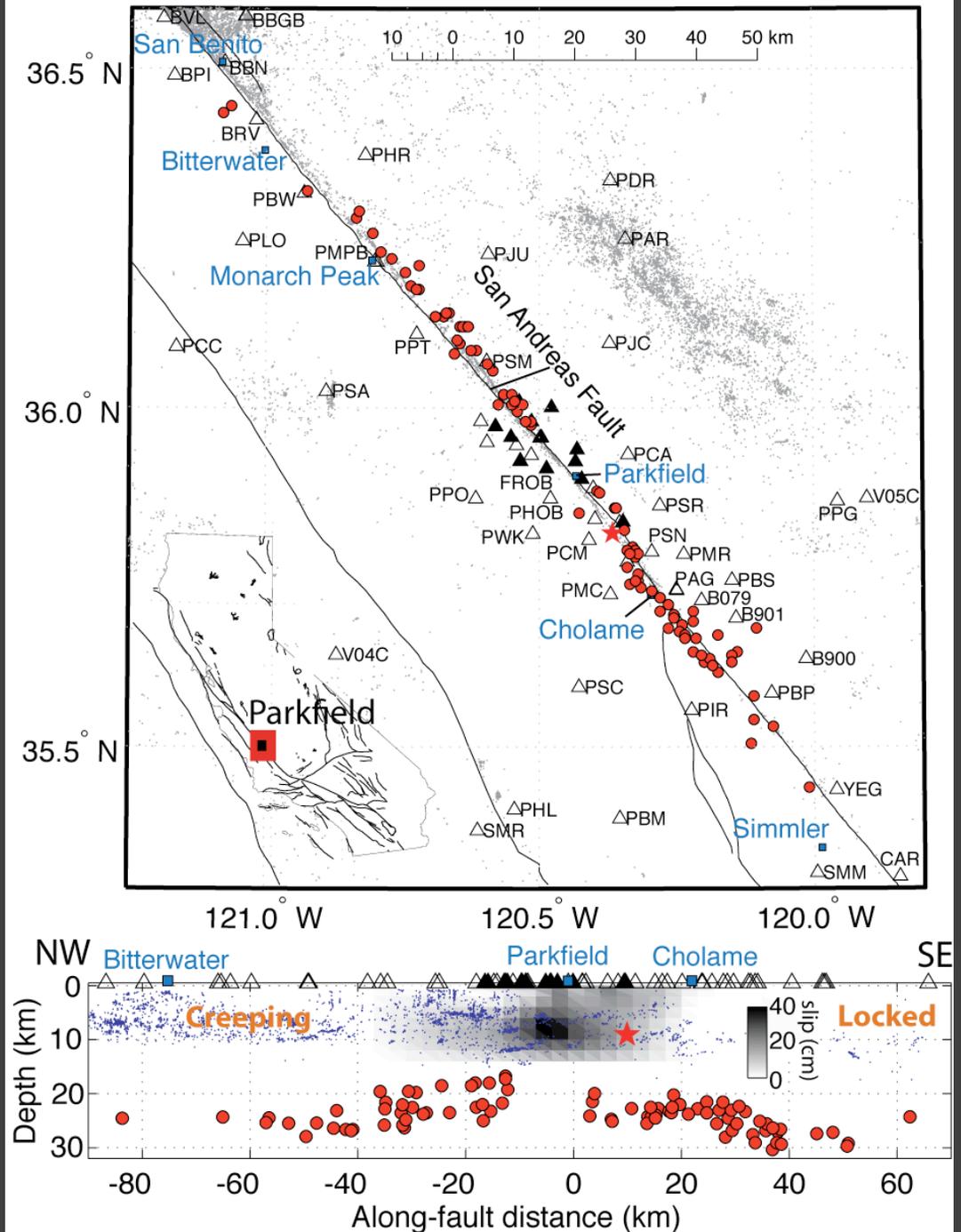


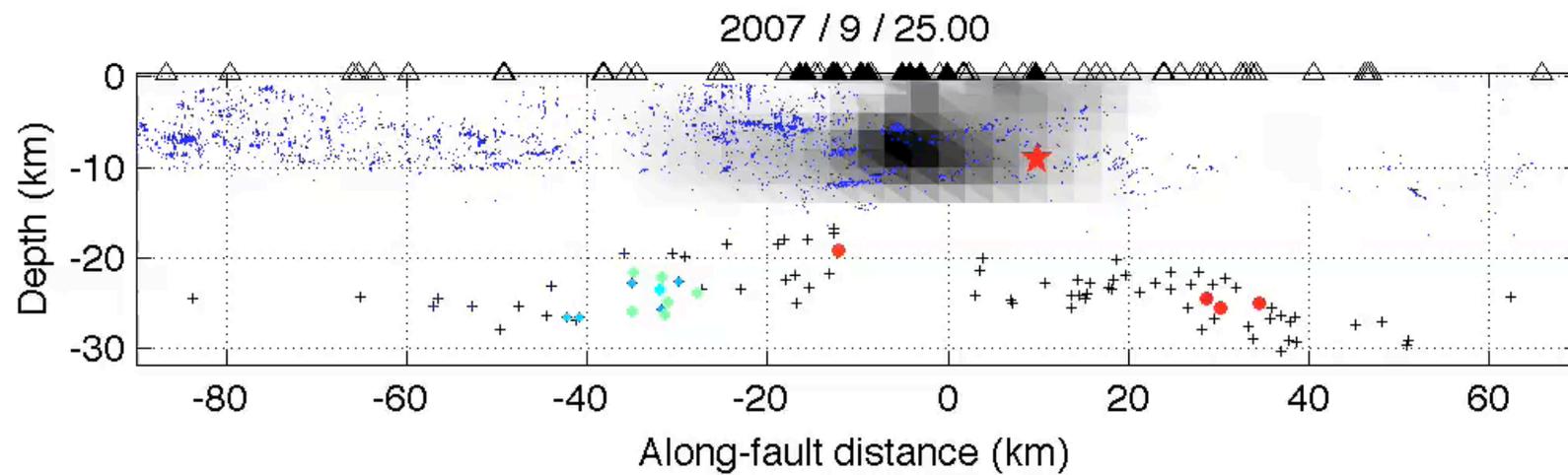
*Shelly and Hardebeck, GRL, 2010*

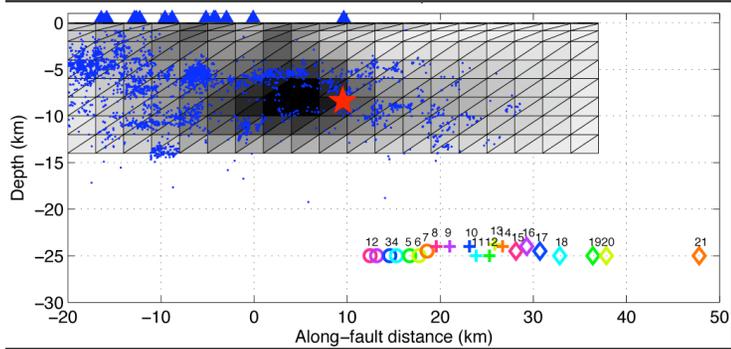
# Parkfield Tremor Locations

- 88 stacked LFE templates
- Located by P and S arrivals on stacked waveforms, using a 3D velocity model.
- Sources extend 75 km both NE and SW of Parkfield

*Shelly and Hardebeck, GRL, 2010*

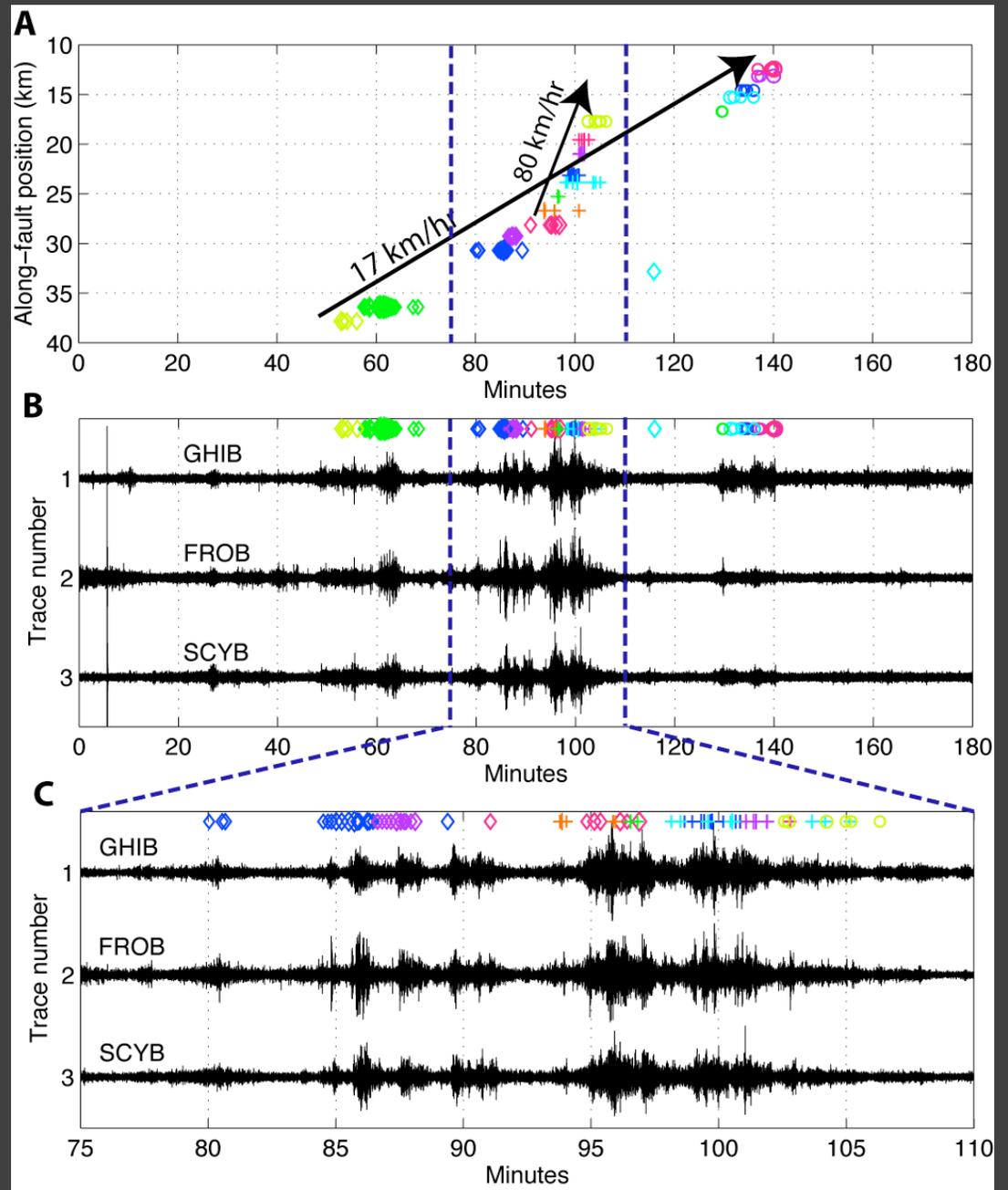






# Tremor Migration

*Shelly, Nature, 2010*



# What does it mean? (1)

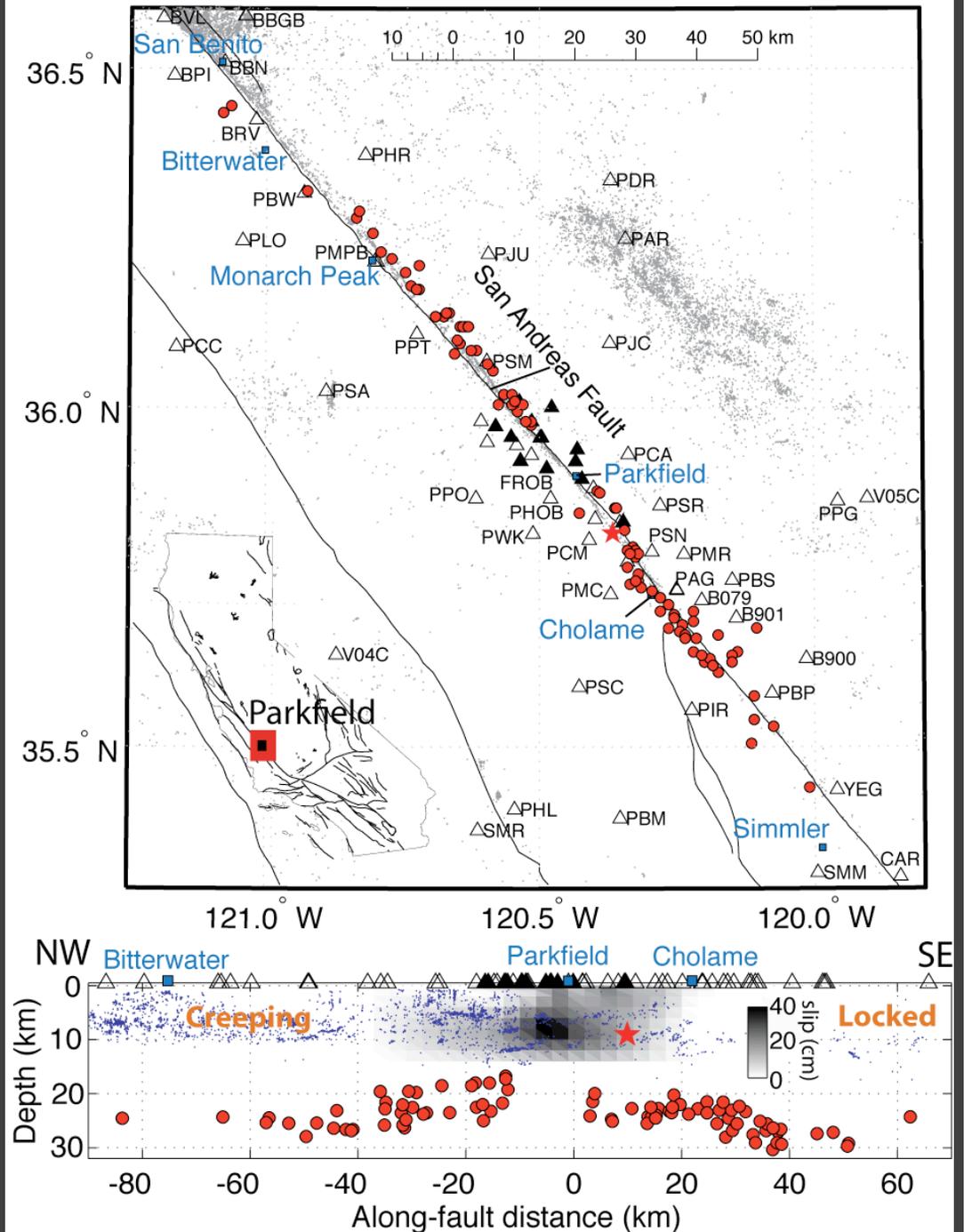
**Conclusion 1.1:** The San Andreas fault does not end at the base of the "seismogenic zone." Tremor sources are located on the deep extension of the fault, in the lower crust. Migration suggests the fault exists as a through-going structure at this depth.

**Conclusion 1.2:** At least some portions of the deep fault deform brittlely. Tremor contains seismic waves of 30+ Hz even with temperatures ~500-600C

# Parkfield Tremor Catalog

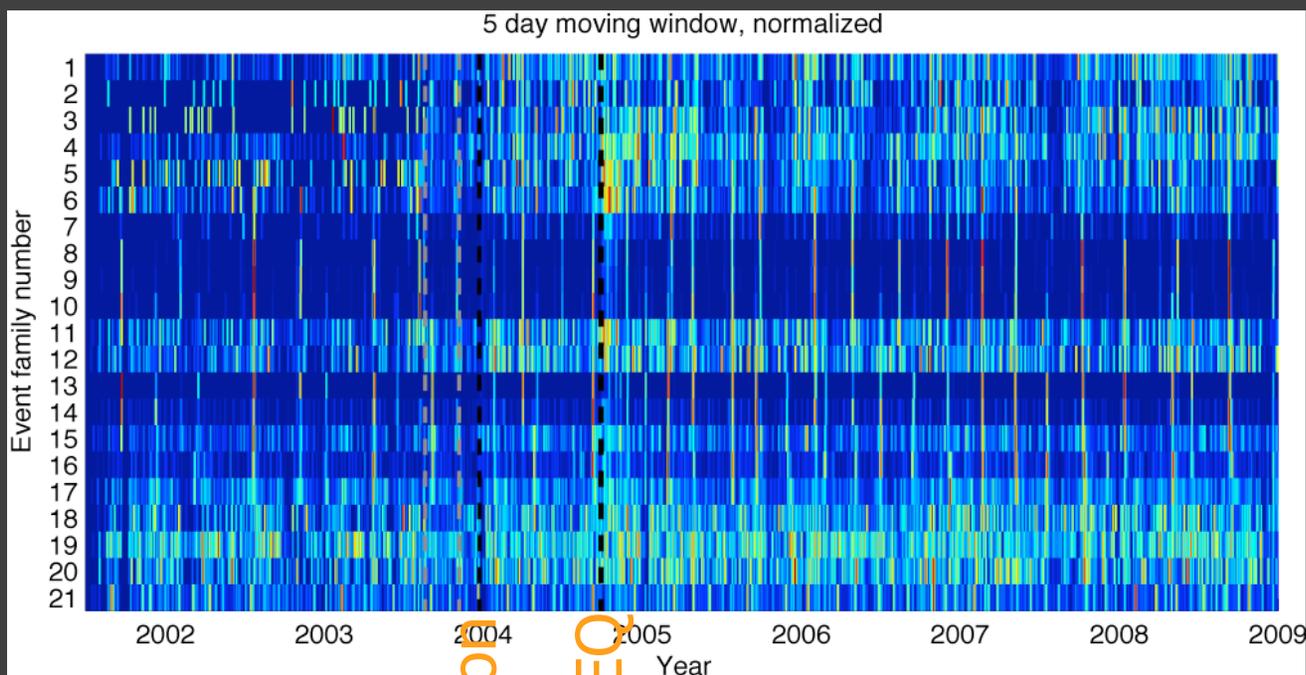
- 9 years of data
- ~36 trillion cross-correlation measurements
- ~600,000 events detected since mid-2001 (3000-20,000 per family)
- Detectable tremor activity in some area every day

*Shelly and Hardebeck, GRL, 2010*

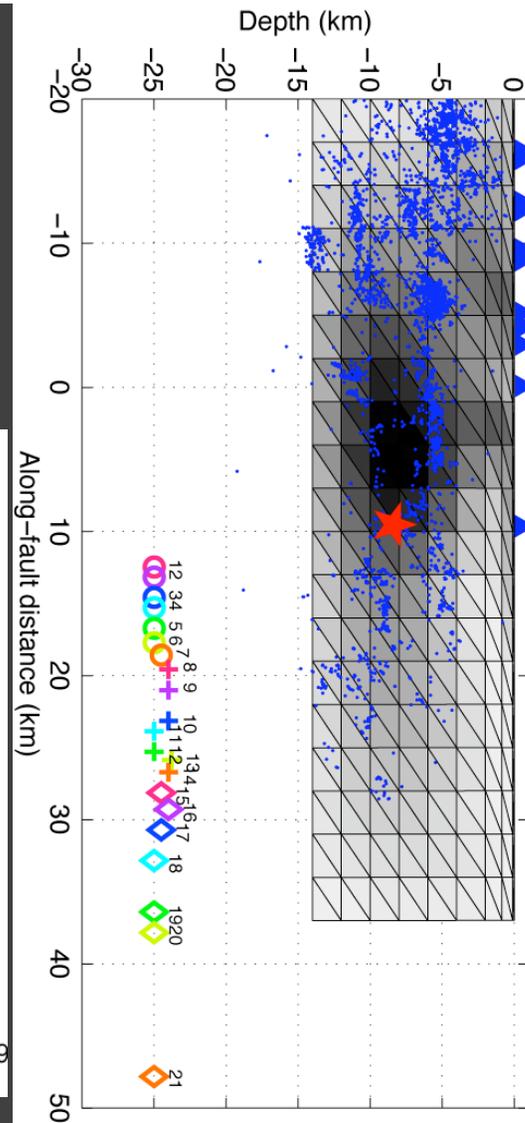


## 2. Variations in recurrence patterns and amplitudes among tremor families

# Event Recurrence Patterns

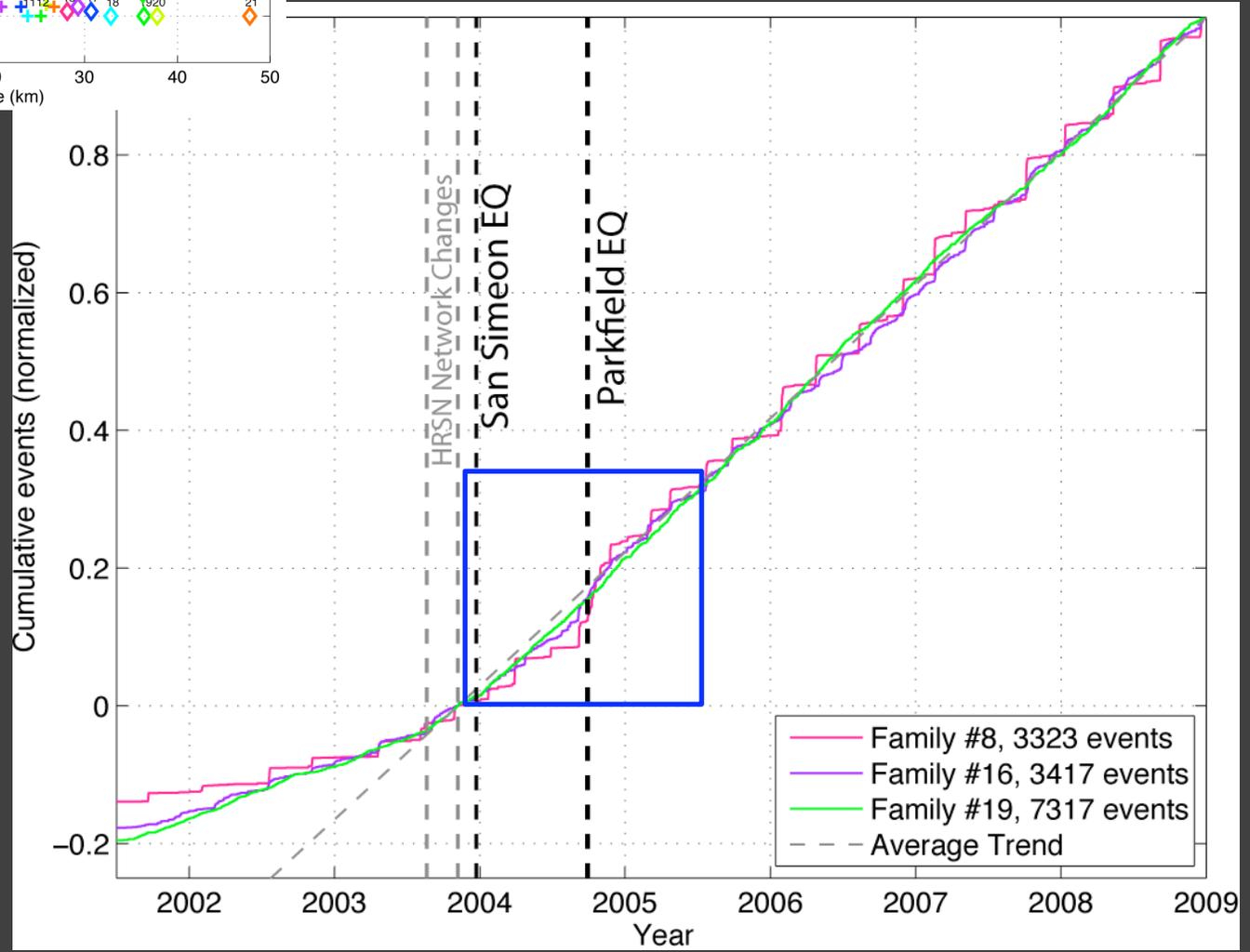
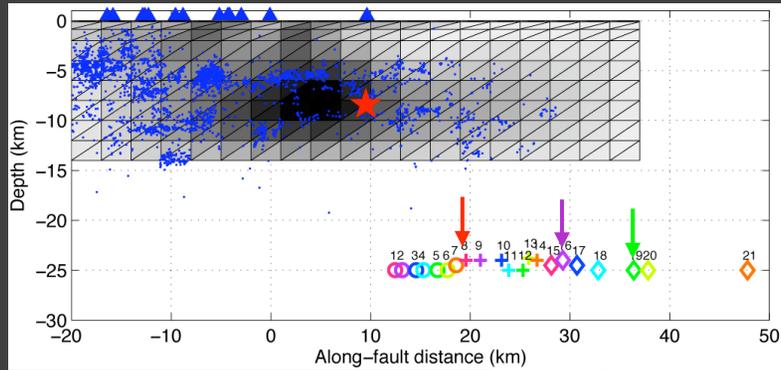


San Simeon  
Parkfield EQ



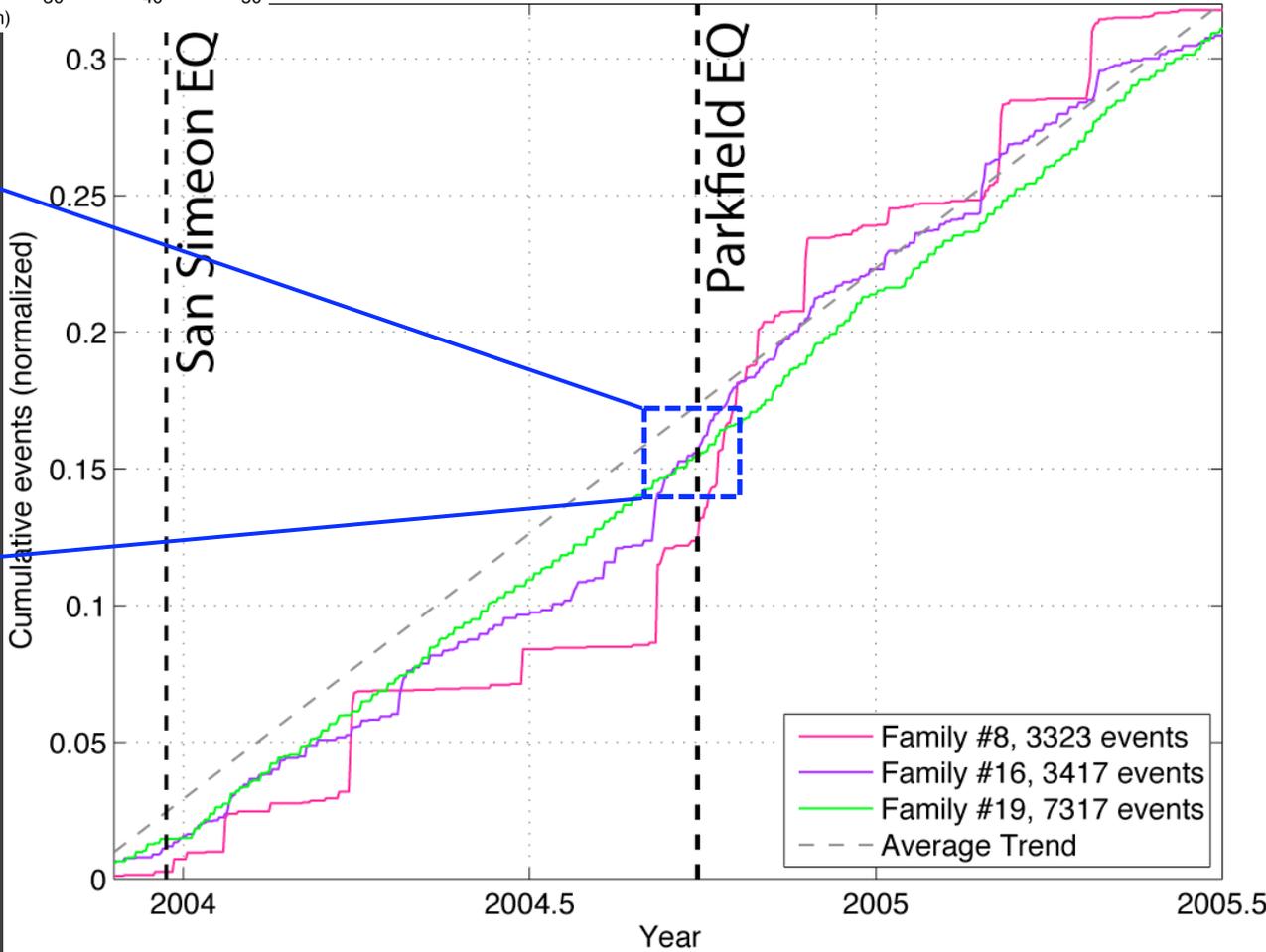
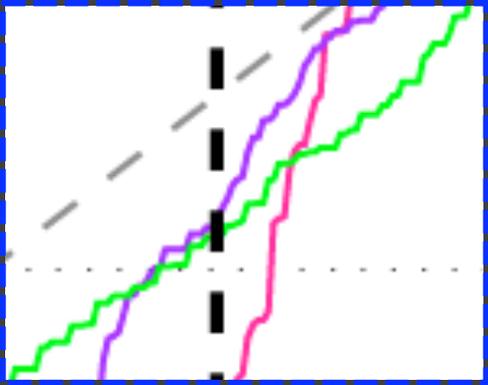
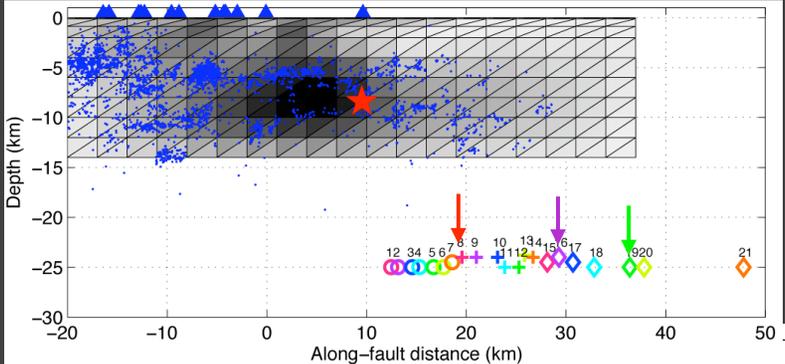
Shelly, Nature, 2010

# Cumulative Events



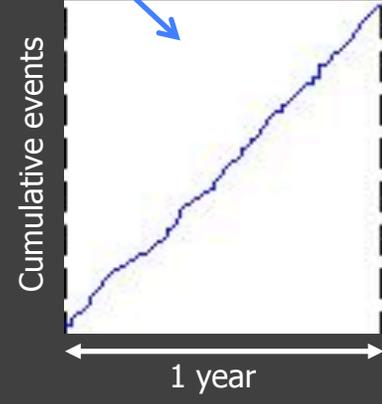
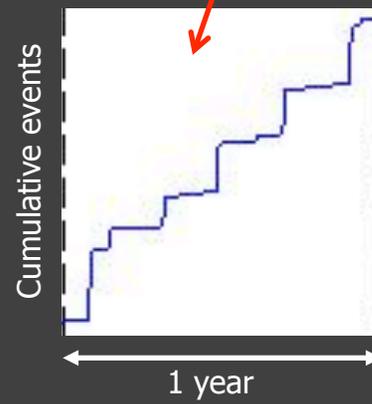
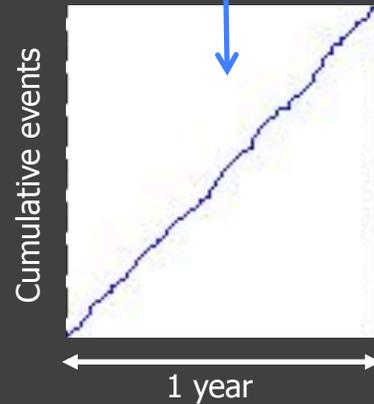
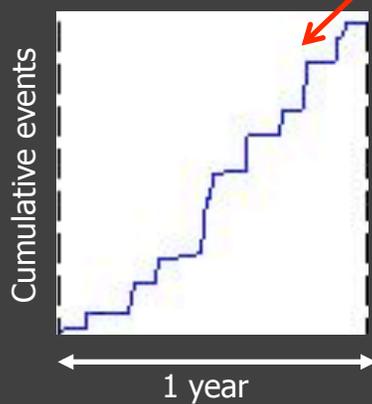
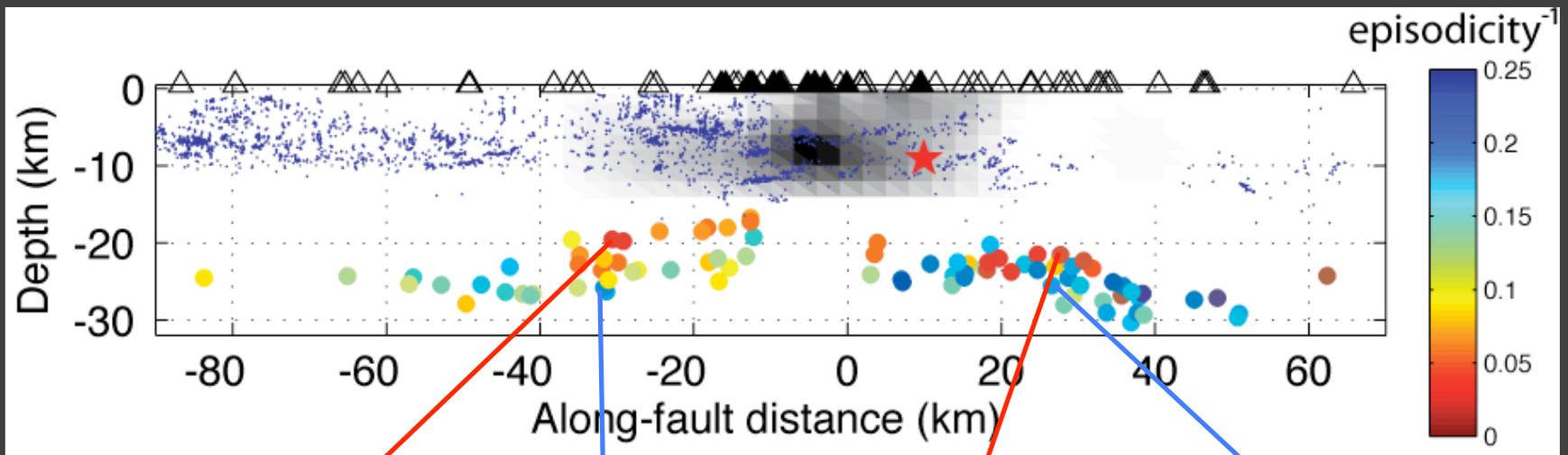
Shelly, Nature, 2010

# Cumulative Events



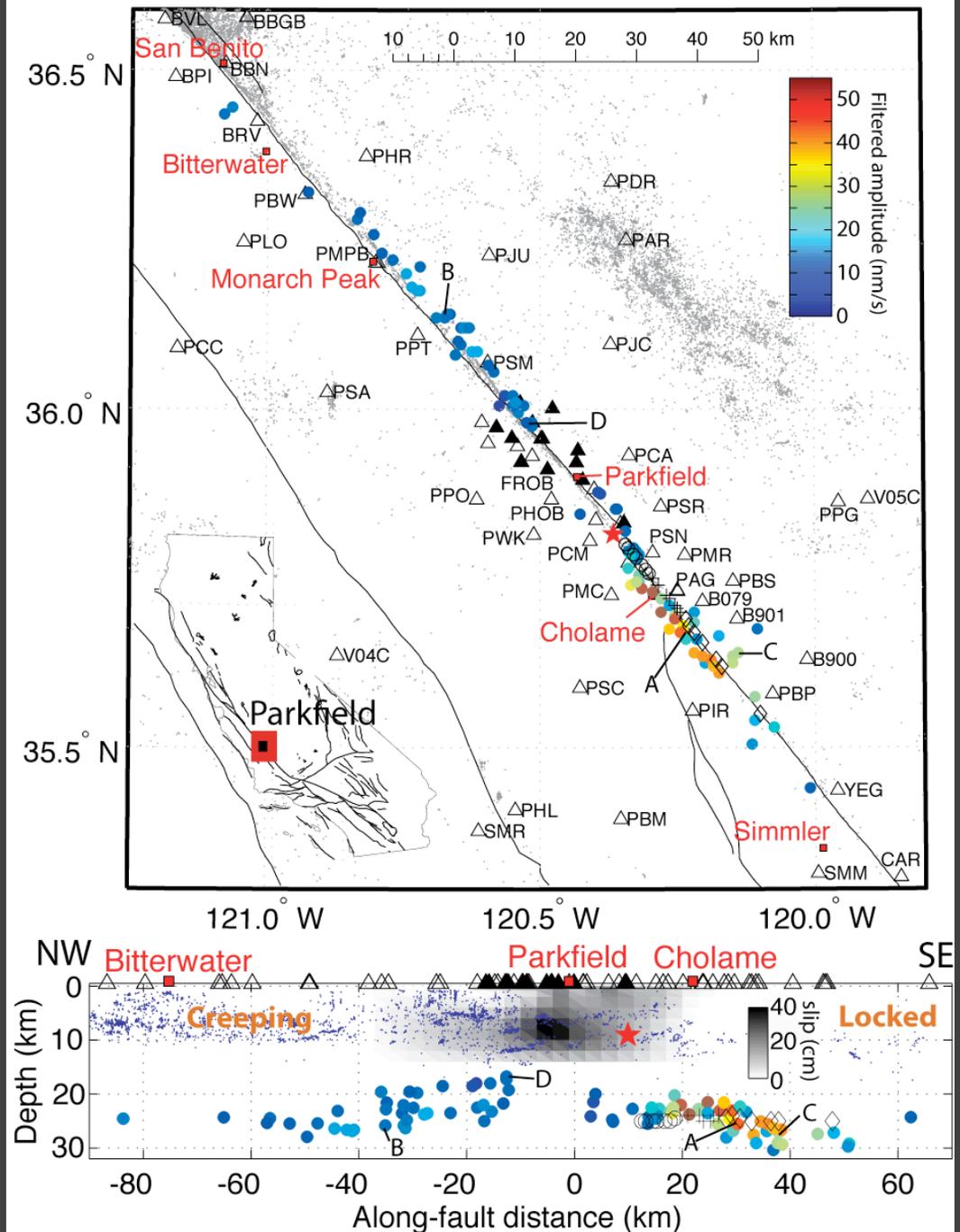
Shelly, Nature, 2010

# Shallower sources have larger, less frequent bursts



# Amplitude potential

- Characterize source amplitude as peak ground velocity of 20<sup>th</sup> largest event during 2001-2010.
- Avoids bias from large amplitude outliers (EQs/ noise) and large number of small amplitude events



# What does it mean? (2)

**Conclusion 2.1:** The strength of the lower crust may vary with depth. Shallower tremor sources have larger, less frequent episodes compared to deeper sources. (But what's happening from 13-20 km depth?)

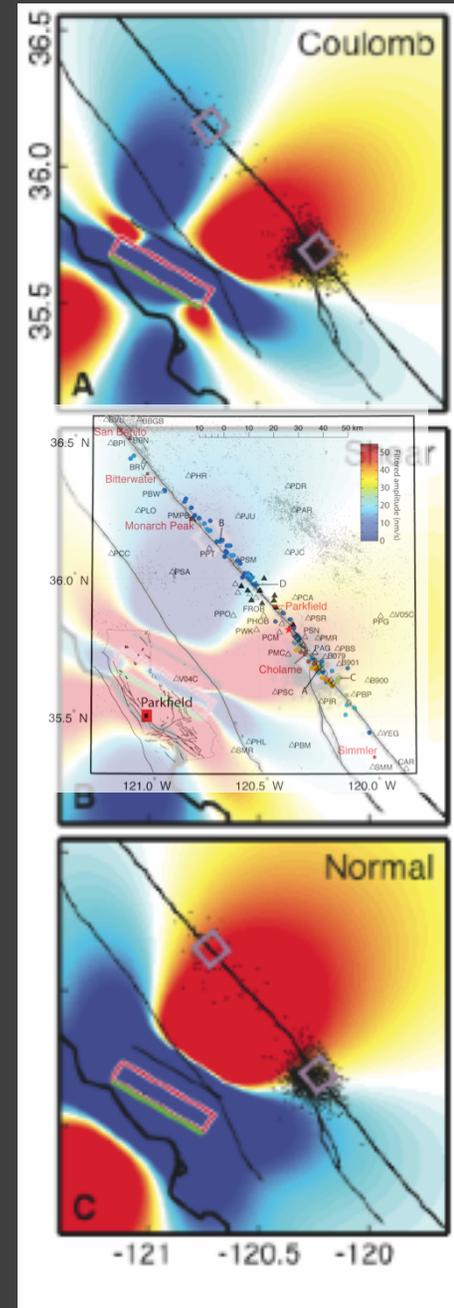
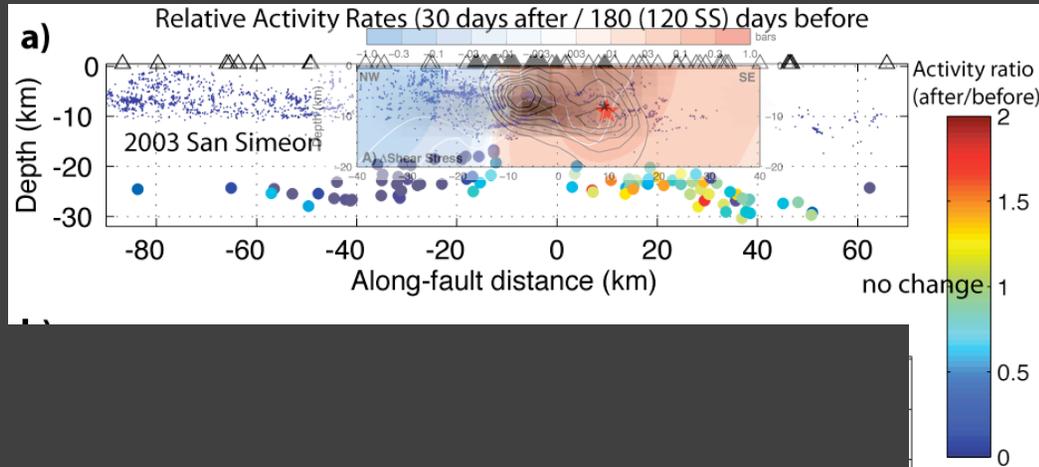
**Conclusion 2.2:** Tremor amplitude varies coherently along strike. This implies a corresponding variation in geology (fluids??). Gap beneath Parkfield may reflect further amplitude variation.

# 3. Tremor response to nearby earthquakes

- 1) 2003 San Simeon (M 6.5)
- 2) 2004 Parkfield (M 6.0)

# Coseismic Stresses

2003 M 6.5 San Simeon  
(after/before activity rates)

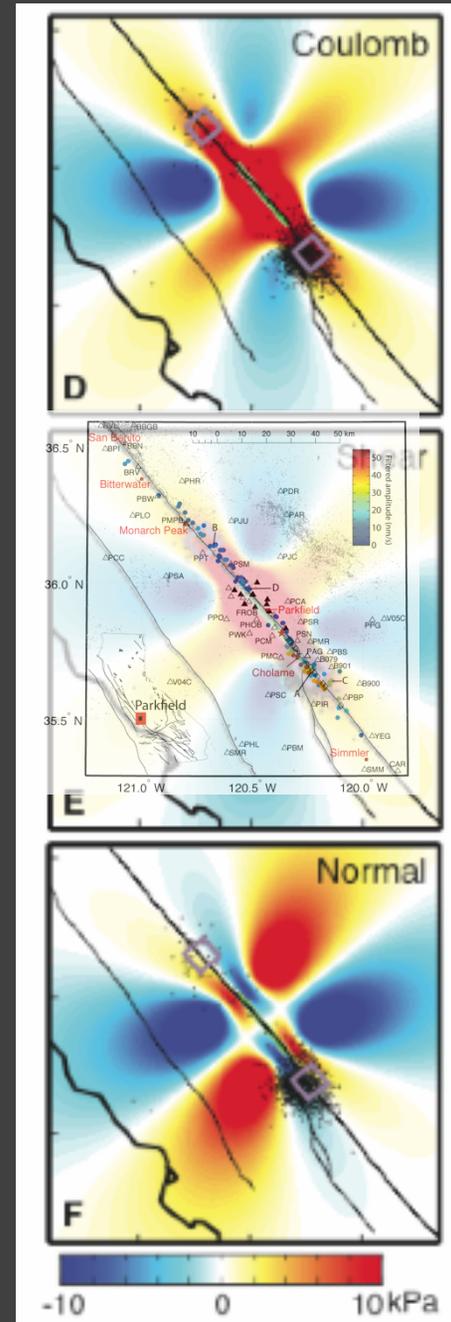
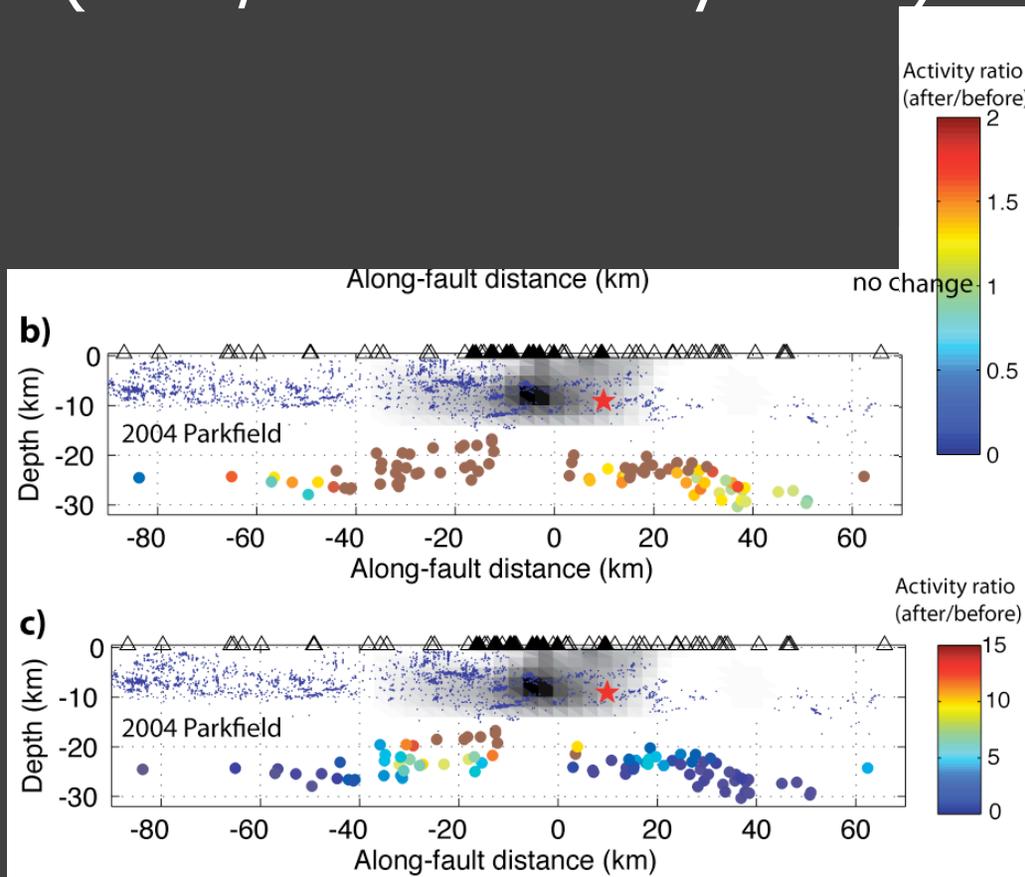


Stress model from Johanson and Bürgmann, 2010

Nadeau and Guilhem, Science, 2009

# Coseismic Stresses

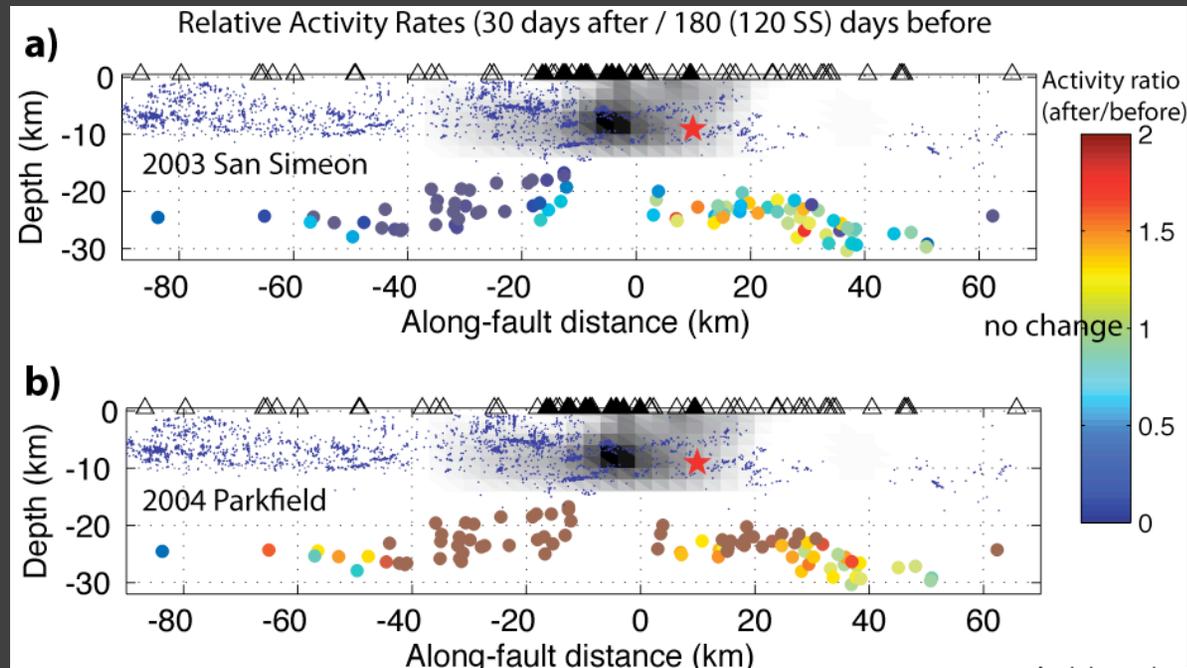
2004 M 6.0 Parkfield  
(after/before activity rates)



Nadeau and Guilhem, *Science*, 2009

# Response to 2003 San Simeon and 2004 Parkfield Earthquakes

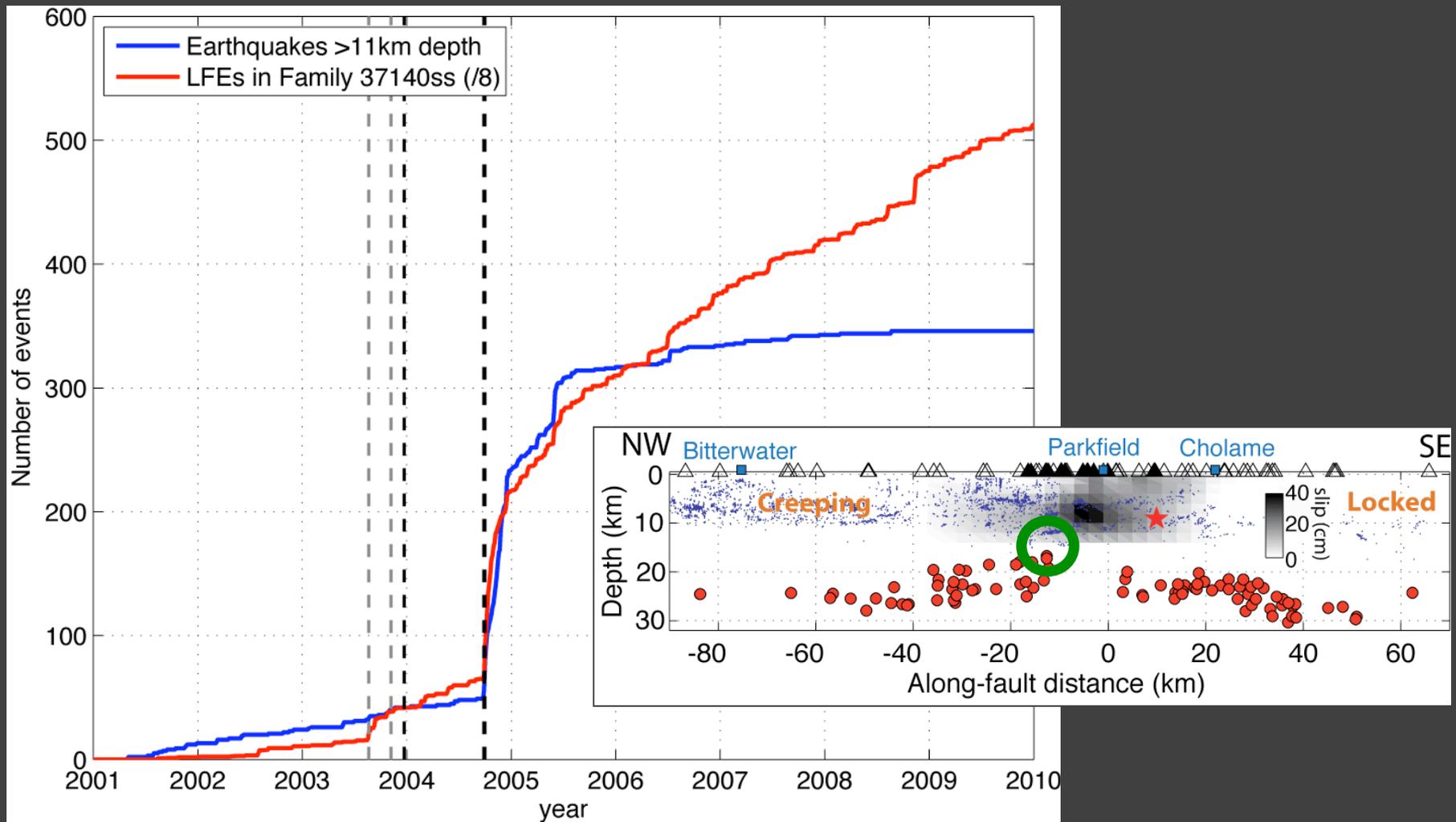
2003 San Simeon



2004 Parkfield

2004 Parkfield  
(rescaled)

# Comparison of Tremor and EQ response



Shelly, in prep.

# What does it mean? (3)

**Conclusion 3.1:** The 2003 San Simeon earthquake produces a strong “stress shadow” effect for the northern tremor sources.

**Conclusion 3.2:** Postseismic deformation following the 2004 Parkfield earthquake extends into the lower crust, probably as deep afterslip. Response is asymmetric, with a greater effect beneath creeping section NW of Parkfield.

# Remaining Questions (Lots!)

1. How do you get brittle (seismic) deformation at 600°C? (*Extreme weakening of the fault???*)
2. Why does the deep fault slip in lots of little tiny events rather than a single larger event? What controls the migration velocities? (*Interplay between brittle and ductile deformation???*)
3. What's happening in places between earthquakes and tremors (~13-20 km depth)? Does this zone slip every few months along with shallow tremor bursts?? Only in big (1857-type) earthquakes???
4. Many more...

