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Meeting Goals:

- Present latest results from EarthScope science
- Emphasize integrated, multi-disciplinary science
- Enhance role of young scientists
- Look to the future

Organizing Committee:

Wayne Thatcher, U.S. Geological Survey Roland Burgmann, University of California, Berkeley Jonathan Price, University of Nevada, Reno

Program Committee:

Aaron Velasco, University of Texas at El Paso Ben van der Pluijm, University of Michigan Peter Shearer, University of California, San Diego Emily Brodsky, University of California, Santa Cruz George Hilley, Stanford University



The EarthScope National Meeting is supported by NSF. Additional funding is provided by NASA and USGS.



Tuesday, March 27

7:00am - 6:00pm	Registration	ELEVATOR LOBBY
8:00am - 1:30pm	Field Trip (pre-registration required) David Schwartz · USGS	Meet in the Portola Plaza Lobby.
8:00am - 9:30pm	Self-Supported Workshops (registration required)	
8:00am - 3:00pm	SNARF Working Group Geoff Blewitt · Nevada Bureau of Mines and Geology	DRIFTWOOD
8:00am - 3:00pm	Preparing for SAFOD Phase 3: Locating the Target Earthquakes and Overview of Core Handling and Analysis Plan Steve Hickman, Bill Ellsworth · U.S. Geological Survey Mark Zoback, Charlie Weiland · Stanford University	COTTONWOOD
8:00am - 3:00pm	MARGINS/GeoSwath Multidisciplinary Workshop: Integrated Collaborations in Cascadia and the Walker Lane/Salton Trough Geoff Abers · Boston University Basil Tikoff · University of Wisconsin	REDWOOD
8:00am - 3:00pm	CIG Training Session for Modeling of Long-Time Scale Geodynamics Michael Gurnis · California Institute of Technology	KAUFMAN
8:00am - 12:00pm	Introduction to Selected DMC Data Access Tools Tim Ahern, Chad Trabant · IRIS	IRONWOOD I
12:30pm - 4:00pm	What Data Products Does PBO Provide and How Do I Get Them? Greg Anderson, Kathleen Hodgkinson · UNAVCO	IRONWOOD II
3:00pm - 7:00pm	Using EarthScope Data in the Classroom John Taber · IRIS Susan Eriksson · UNAVCO	REDWOOD I
3:00pm - 6:00pm	Continental Intraplate Deformation & Seismicity: What We Know, What We Don't, and What We Need To Seth Stein · Northwestern University	IRONWOOD
7:30pm - 9:30pm	Xenoliths & EarthScope Workshop Randy Keller · University of Oklahama	IRONWOOD

3:00pm – 6:30pm	Mini-Courses	
3:00pm - 3:45pm	Geochemistry and Petrology of the Crust and Mantle Made Simple Mihai Ducea · University of Arizona	BONSAI BALLROOM
3:45pm - 4:30pm	Tomography/Seismic Imaging Suzan van der Lee · Northwestern University	BONSAI BALLROOM
4:30pm – 5:00pm	Break	
5:00pm - 5:45pm	How Anyone Can Access EarthScope Data Greg Anderson (PBO) · UNAVCO Chad Trabant (USArray) · IRIS Consortium Charlie Weiland (SAFOD) · Stanford University	BONSAI BALLROOM
5:45pm - 6:30pm	GPS/InSAR and EarthScope Matt Pritchard · Cornell University	BONSAI BALLROOM

Wednesday, March 28

6:00am - 6:00pm	Registration	PORTOLA LOBBY
7:00am – 7:45am	Breakfast	DE ANZA BALLROOM
8:00am - 10:00am	Plenary Session: Building and Destroying a Volcano in Real Time Michael Lisowski · U.S. Geological Survey Stephanie Prejean · Alaska Volcano Observatory	STEINBECK FORUM
10:00am – 10:30am	Break	DE ANZA FOYER
10:30am - 11:15am	New Opportunities for EarthScope Science: Perspectives from NSF, NASA & USGS Kaye Shedlock · NSF John LaBrecque · NASA Linda Gundersen · U.S. Geological Survey	STEINBECK FORUM
11:15am - 12:00pm	Poster Session: Open	SERRA BALLROOM
12:00pm – 1:00pm	Lunch	DE ANZA BALLROOM
1:00pm – 3:00pm	Plenary Session: What Have We Learned About Faults and Earthquakes? Yuri Fialko · Scripps Institution of Oceanography John Solum · U.S. Geological Survey	STEINBECK FORUM
3:00pm – 3:30pm	Break	DE ANZA FOYER
3:30pm - 5:30pm	Plenary Session: Non-Volcanic Tremor and Episodic Slip: Fundamental Processes or Curiosities? David Shelly · Stanford University	STEINBECK FORUM
5:30pm - 7:00pm	Poster Session: -Building and Destroying a Volcano in Real Time -What Have We Learned About Faults and Earthquakes? -Non-Volcanic Tremor and Episodic Slip: Fundamental Processes or Curiosities?	SERRA BALLROOM
7:30pm - 9:30pm	Reception with hors d'oeuvres and Speaker at the Monterey Bay Aquarium Extending Seismic Networks to the Deep Sea Marcia McNutt · Monterey Bay Aquarium Research Institute	Meet in the Portola Plaza Lobby. Shuttle leaves starting at 7:00pm

Thursday, March 29

6:00am - 6:00pm	Registration	PORTOLA LOBBY
7:00am – 7:45am	Breakfast	DE ANZA BALLROOM
8:00am - 10:00am	Plenary Session: Mantle Lithosphere: Crucial or Irrelevant to Continental Deformation? Hersh Gilbert, Purdue University William Holt · State University of New York, Stony Brook	STEINBECK FORUM
10:00am – 10:30am	Break	DE ANZA FOYER
10:30am - 12:30pm	Plenary Session: Bringing EarthScope Research into the Undergraduate Classroom Barbara Tewksbury · Hamilton College	STEINBECK FORUM
12:30pm – 1:30pm	Lunch	DE ANZA BALLROOM
1:30pm - 3:15pm	Poster Session: -Mantle Lithosphere: Crucial or Irrelevant to Continental Deformation? -The New Madrid Seismic Zone -Quality Control, Data Centers and New Tools -New Technology and Methods	SERRA BALLROOM
3:15pm – 3:45pm	Break	DE ANZA FOYER
4:00pm - 6:00pm	Plenary Session: What Drives the Western United States and How Does it Deform? Lucy Flesch · Purdue University	STEINBECK FORUM
6:00pm – 7:00pm	Dinner on your own	
7:00pm - 9:00pm	Poster Session: -What Drives the Western US & How Does it Deform? -Building a Continent in 4D -Infrastructure	SERRA BALLROOM

Friday, March 30

7:00am – 7:45am	Breakfast	LOWER ATRIUM	
8:00am - 10:00am	Plenary Session: Building a Continent in 4D Basil Tikoff · University of Wisconsin, Madison	STEINBECK FORUM	
10:00am – 10:30am	30am Break		
	Closing Plenary Session: Future Directions in EarthScope Science		
	EarthScope and the Integration of Synoptic Data: "Understanding the Structure and Evolution of a Continent Anne Meltzer · Lehigh University		
10:30am - 12:30pm	Reconciling Deformation and Rheology Using EarthScope Data Thomas Herring · Massachusetts Institute of Technology	STEINBECK FORUM	
	Broad Band Seismology at the Blue End of the Spectrum: Opportunities and Challenges for EarthScope Science William Ellsworth · U.S. Geological Survey		



• 5

I. Building and Destroying a Volcano in Real Time

Co-Chairs: Wayne Thatcher, U.S. Geological Survey; Aaron Velasco, University of Texas at El Paso

Monitoring crustal deformation and seismic activity and imaging subsurface structure of magmatic systems provides fundamental constraints on how volcanoes grow and evolve. EarthScope monitoring at Mt. St. Helens and Augustine is already supplying information on how these processes occur in near real time, providing a snapshot on how these volcanoes are built and destroyed.

8:00am Invited Plenary Talks

Under Construction: What We've Learned About Volcano Building from the Ongoing Eruption at Mount St. Helens · Michael Lisowski et al. Tracking Magma Ascent in the 2006 Eruption of Augustine Volcano, Alaska: The Role of EarthScope · Stephanie Prejean et al.

9:00am Short Contributed Talks

Time-Dependent Finite Element Modeling of Effusive Silicic Eruptions · Kyle Anderson and Paul Segall

Accelerated Uplift of the Yellowstone Caldera, 2004-2006, From GPS and InSAR Observations · Wu-Lang Chang et al.

Eruption Dynamics at Mount St. Helens Imaged from Inversion of Broadband Seismic Wafeforms: Interaction of the Shallow Magmatic and Hydrothermal Systems · Gregory Waite et al.

9:30am Panel Discussion

Michael Lisowski, Stephanie Prejean, Paul Segall, David Hill

II. What Have We Learned About Faults and Earthquakes?

Co-Chairs: Emily Brodsky, University of California, Santa Cruz; Ben van der Pluijm, University of Michigan

A major goal of EarthScope is to figure out why earthquakes happen. What are the mechanics of starting, propagating and stopping a sudden slip event? The SAFOD project directly targets the earthquake source zone through a variety of geophysical and geological tools focused on the San Andreas fault. Other parts of EarthScope are designed to capture the longer term geodetic deformation associated with earthquakes or the shorter term shaking. This session will high-light the results with an emphasis on combining the insights from multiple tools.

1:00pm Invited Plenary Talks

Space Geodetic Imaging of Deformation Due to Active Faults Throughout the Earthquake Cycle · Yuri Fialko What Have we Learned About Faults and Earthquakes? Using a View from SAFOD to Increase Understanding of Fault Behavior · John Solum

2:15pm Short Contributed Talks

Contributions of EarthScope to Earthquake Hazard Assessment in Northern California · Tom Brocher et al.

Early Insights into the Mechanical Behavior of Materials in the 3D SAFOD Volume · Brett Carpenter et al.

Similarity of Paleostress and in Situ Stress at SAFOD and Implications for a weak San Andreas Fault · Rafael Almeida et al.

Kinematic Modeling of the SAFOD Target Events · Douglas Dreger et al.

III. Non-Volcanic Tremor and Episodic Slip: Fundamental Processes or Curiosities?

Co-Chairs: Roland Burgmann, University of California, Berkeley; Emily Brodsky, University of California, Santa Cruz

EarthScope is focusing significant effort on the two regions where non-volcanic tremor and/or slow slip events have been recognized, including the deployment of borehole strainmeters, seismometers and GPS stations along the Cascadia subduction zone and in the Parkfield/Cholame region. This session will present fundamental issues about the nature and implications of observed transient deformation events and seismic tremor activity, which have been recognized along some subduction zones as well as strike-slip faults in Japan and California. Not all slow deformation events have known tremor activity associated with them and not all tremor events have been accompanied by resolvable deformation transients. What is needed to clarify what processes (slip, fluid flow, and/or others) are associated with the observed phenomena (tremor, transient surface deformation, long-period earthquakes, triggered seismicity)? What are the implications of these transient processes for earthquake occurrence and hazard estimates?

3:30pm Invited Plenary Talk

The Mechanics of Episodic Tremor and Slip · David Shelly et al.

4:20pm Short Contributed Talks

A 70+ Station GPS Recording of the January, 2007 Cascadia ETS · Tim Melbourne Slow-Slip and Triggered Earthquakees on Kilauea Volcano with Implications for Slow-Slip and Tremor in Subduction Zones · Paul Segall et al. Do Episodic Tremor and Slip (ETS) Events Affect Seismicity in the Northern Cascadia Subduction Zone? · Tom Pratt Non-Volcanic Tremor Driven by Large Transient Shear Stresses · Justin Rubinstein et al. Imaging Subduction, Episodic Tremor and Slip in the Pacific Northwest: Cascadia Arrays for EarthScope (CAFE) · Geoffrey Abers et al. Tremor · Aaron Wech and Kenneth Creager Effect of Parkfield Earthquake on Tremor Activity Below the San Andreas Fault Near Cholame, CA · Bob Nadeau

IV. Mantle Lithosphere: Crucial or Irrelevant to Continental Deformation?

Co-Chairs: Peter Shearer, University of California, San Diego; Aaron Velasco, University of Texas at El Paso

There is long-standing debate over the role of the mantle in the development of topographic features, plate movement, deformation zones, and continental formation. Questions include the role of a weak asthenosphere and ductile flow directions associated with the orogenies, the importance of delamination and upper-mantle buoyancy forces, and the degree of coupling among plates, microplates and the mantle flow field. Data currently being collected and analyzed by EarthScope will help to resolve these issues through detailed mapping of surface deformation, seismic observations of mantle structures, and geodynamic modeling of stress and mantle flow. This research includes direct geodetic observations of the surface strain field, seismic imaging of velocity anomalies, correlation of surface geologic features with seismic structures at depth, shear-wave splitting constraints on the mantle flow field, and modeling of plate boundary dynamics and driving forces for continental deformation. This session explores geodetic, seismic, and geological evidence for the role of mantle dynamics on continental formation and deformation.

8:00am Invited Plenary Talks

Lithosphere Coupling with Mantle Circulation in Western North America, Central Asia, and the Rest of the World · William Holt et al. The Sierra Nevada: Evidence of the Importance of Mantle Lithosphere in Continental Deformation · Hersh Gilbert et al.

9:00am Short Contributed Talks

Lithosphere Temperature, Strength, and Deformation in Western North America and EarthScope · Roy Hyndman et al.

Colorado Rockies Experiment and Seismic Transects (CREST): Cenozoic Uplift, Magmatism, and Mantle to Surface Fluid Interconnections Associated with the Aspen Anomaly · Karl Karlstrom et al.

Mapping Upper-Mantle Anisotropy Beneath the Western US: Toward a Coupled Seismic and Geodynamic Analysis of Crust-Mantle Coupling -James Gaherty et al.

9:30am Panel Discussion

William Holt, Hersh Gilbert, Roy Hyndman, James Gaherty, Karl Karlstrom

V. Bringing EarthScope Research into the Ungergraduate Classroom

Co-Chairs: Michael Hamburger, Indiana University; Aaron Velasco, University of Texas at El Paso

The EarthScope initiative provides unprecedented opportunities to both the research and education communities. As participants in the EarthScope experiment, many of us bring our experience and expertise with state-of-the-art geophysical and geological data into the classes we teach. This session will allow participants to share their experiences—ranging from tried-and-true exercises to fresh, untested new ideas—on the use of EarthScope-related data and results into the undergraduate science classroom. We invite participants to contribute a poster focused on educational applications of EarthScope science, and to bring along a lab exercise, homework problem, or new idea for using EarthScope data in the classroom.

10:30am	Invited Plenary Talk From Data to Insight: Developing Effective Undergraduate Activities Using EarthScope Data · Barb Tewksbury
11:00am	Break Out/Discussion EarthScope Data in the Classroom · Michael Hamburger, Indiana University; Barb Tewksbury, Hamilton College
11:20am	Poster Summaries
11:40am	Poster Session

VI. What Drives the Western US & How Does it Deform?

Bringing EarthScope Research into the Undergraduate Classroom

Co-Chairs: George Hilley, Stanford University; Wayne Thatcher, U.S. Geological Survey

The kinematics of the Western US represents the balance between plate-boundary loading, normal and shear tractions acting along the base of the lithosphere, internal lithospheric buoyancy, local surface processes, and lithospheric rheology. USArray will provide constraints on the structure of the lithosphere that may reflect lateral and vertical rheological variations and constrain lithospheric buoyancy. Plate Boundary Observatory (PBO) will provide a clear picture of the deformation-rate field over geodetic (i.e., decadal) time scales. Over these geodetic time scales, the deformation-rate field records processes related to the long-term loading of the plate margins, interseismic earthquake cycle effects, and near-surface hydrologic processes. Geologic studies that constrain the kinematics of the Western US over thousands to millions of years lack the spatial resolution of geodetic instrumentation, but contain information about changes in plate boundary loading and rheology, while filtering out many of the shorter time-scale processes (e.g., hydrologic signals and earthquake-cycle effects) that may confound inferences of plate-boundary loading. Thus, the analysis of the deformation-rate field in the Western US over various time scales should provide important clues to the relative magnitude of the driving forces that produce the deformation, and how these forces have changed over time. In this session, we invite posters and short contributions that use geodetic and/or geologic data to quantify the kinematics of the Western US. In addition, geodynamic studies that seek to explain such observations in terms of the driving forces and rheology of the Western US are also encouraged. We especially seek contributions that link specific observations of the kinematics of the Western US over various time scales to its dynamics through geodynamic models.

4:00pm Invited Plenary Talk

Driving Forces of the Western North American Plate Boundary Zone · Lucy Flesch et al.

4:30pm Short Contributed Talks

Dislocation Models of Interseismic Deformation in the Western United States · Fred Pollitz et al.

- The Walker Lane: How Complex is It? Geodetic and Geologic Strain in the Western Basin and Range Using Enhanced Block Modeling Algorithms · William Hammond et al.
- Rotational Mantle Flow Beneath the Western US · George Zandt and Gene Humphreys
- The Uplift of the Southern Sierra Nevada, the Isabella Seismic Anomaly and Death Valley Extension: A Single Geodynamic Process? · Laetitia Le Pourhiet and Michael Gurnis

VII. Building a Continent in 4D

Co-Chairs: Ben van der Pluijm, University of Michigan; Jonathan Price, University of Nevada, Reno

As a major geophysical experiment, EarthScope is providing heretofore unprecedented views of the structure of the North American continent today. The geodetic observations, and to some extent the seismic observations, will give us a short-term view of geological changes (signals in time frames from seconds to several years), but full four-dimensional analysis, back into deep time, will require the integration of EarthScope's geophysical observations with geochronology, structural geology, neotectonics, mineral physics, petrology, geochemistry, and detailed geologic mapping. This session will highlight some of the promising areas for this integration.

8:00am Invited Plenary Talk

Building a Continent: The Scientific Motivation Behind a Coast-to-Coast GeoSwath · Basil Tikoff

8:40am Short Contributed Talks

Processes of Terrarne Accretion and Modification: The Klamath Mountains of Cascadia · Kate Miller and Calvin Barnes

Regional Geophysical Analyses to Patch Together the Pieces of the GeoSwath · Randy Keller et al.

EarthScope in 'Fly-Over' Country . . . Exploring the Structure and Geological Evolution of the Crust and Lithosphere of the Continental Interior -Ernest Hauser

A Continent-Wide 1-Hz Map of Lg Code Q Variation Across Eurasia and its Relation to Lithospheric Evolution · Brian Mitchell et al.

9:20am Panel Discussion

A

Lon Abbott	Red Rocks Community College
Geoffrey Abers	Boston University
Duncan Agnew	University of California, San Diego
Tim Ahern	IRIS
Sheyla Alayón	University of Puerto Rico at Mayaguez
Richard Allen	UC Berkeley
Rafael Almeida	Texas A & M University
Mark Alvarez	IRIS
Mohamed Aly	Idaho State University
Kent Anderson	IRIS
Greg Anderson	UNAVCO
Kyle Anderson	Stanford University
Victoria Andreatta	UNAVCO, Inc.
Elizabeth Anthony	University of Texas at El Paso
Edwin Apel	University of California, Berkeley
David Applegate	U.S. Geological Survey
Ramon Arrowsmith	Arizona State University
Richard Aster	New Mexico Tech
Luciana Astiz	University of California San Diego
Gary Axen	New Mexico Inst. of Mining and Technology

В

Hassan Babaie	Georgia State University
Chaitan Baru	University of California, San Diego
Andre Basset	UNAVCO inc
Bruce Beaudoin	IRIS PASSCAL
Susan Beck	University of Arizona
Heather Bedle	Northwestern University
Caroline Beghein	Arizona State University
Rick Bennett	University of Arizona
Mark Benthien	Southern California Earthquake Center
Susan Bilek	New Mexico Tech
Terrence Blackburn	Massachusetts Institute of Technology
Brian Blackman	UNAVCO
Michael Blanpied	U.S. Geological Survey
Thomas Bleier	QuakeFinder
Geoffrey Blewitt	University of Nevada, Reno
Karin Block	Lamont-Doherty Earth Observatory
Frederick Blume	UNAVCO
Yehuda Bock	Scripps Institution of Oceanography

Paul Bodin	University of Washington
Mark Brandon	Yale University
Richard Brazier	Penn State University DuBois Campus
Richard Briggs	Caltech
Tom Brocher	U.S. Geological Survey
Emily Brodsky	University of California, Santa Cruz
Daniel Brothers	University of California, San Diego
Justin Brown	Stanford University
Larry Brown	Cornell University
Kevin Brown	Scripps Institution of Oeanography
Scott Burdick	Massachusetts Institute of Technology
Roland Burgmann	University of California, Berkeley
Robert Busby	IRIS
Rhett Butler	IRIS
Katrina Byerly	University of South Carolina
C	
Eric Calais	Purdue University
Warren Caldwell	Stanford University
Josh Calkins	University of Arizona
Keith Cardon	University of Texas at El Paso
Brett Carpenter	Pennsylvania State University
Rufus Catchings	U.S. Geological Survey
Wu-Lung Chang	University of Utah
J. Andres Chavarria	Paulsson Geophysical
Judith Chester	Texas A&M University
Frederick Chester	Texas A&M University
Elizabeth Cochran	University of California, Riverside
Brian Coyle	UNAVCO
Kenneth Creager	University of Washington
Christopher Crosby	Arizona State University
D	
Jeffrey Dingler	University of California, San Diego
Tim Dittman	UNAVCO

Tim Dittman UNAVCO
Mai Linh Doan University of California, Santa Cruz
Andrea Donnellan Jet Propulsion Laboratory
Perle Dorr IRIS
Briget Doyle College of Charleston
Douglas Dreger UC Berkeley
Mihai Ducea University of Arizona
Ken Dueker University of Wyoming

Clark Dunson	QuakeFinder	Priscilla Grew	University of Nebraska-Lincoln
Adam Dziewonski	Harvard University	Aurelie Guilhem	University of California, Berkeley
E		Linda Gundersen	U.S. Geological Survey
Kevin Eagar	Arizona State University	Yonggui Guo	Rice University
Jennifer Eakins	University of California, San Diego	Cansun Guralp	Guralp Systems, Ltd.
Arias Eliana	IRIS/PASSCAL	Michael Gurnis	California Institute of Technology
William Ellsworth	U. S. Geological Survey	н	
Susan Eriksson	UNAVCO	Jafar Hadizadeh	University of Louisville
James Evans	Utah State University	Katrin Hafner	UNAVCO, Inc.
F		Bezalel Haimson	University of Wisconsin
Iamia Farrell	University of Litab	Michael Hamburger	Indiana University
Noah Fav		William Hammond	Nevada Bureau of Mines and Geology
Karl Feaux		Samantha Hansen	University of California, Santa Cruz
Keedan Fendler		Steven Hansen	University of Wyoming
Yuri Fialko		Roger Hansen	University of Alaska Fairbanks
Fric Fielding	Caltech	Steven Harder	University of Texas at El Paso
Melvin Fillerup	University of South Carolina	Rebecca Harrington	University of California, Los Angeles
Delphine Fitzenz	CNRS - Geophysics Institute Strasbourg	Robert Harris	Oregon State University
Lucy Flesch.	Purdue University	Derrick Hasterok	University of Utah
David Forand.	Utah State University	Michael Hasting	UNAVCO Inc
Matthew Fouch	Arizona State University	Ernest Hauser	Wright State University
David Fountain	National Science Foundation	Gavin Hayes	Penn State University
James Fowler	IRIS	John He	GSC Pacific
Andrew Frassetto	University of Arizona	Margaret Hellweg	UC Berkeley
Jeff Freymueler	University of Alaska, Fairbanks	Don Helmberger	Caltech
Gary Fuis	U.S. Geological Survey	Daniel Hernandez	Stony Brook University
Junichi Fukuda	Indiana University	Thomas Herring	Massachusetts Institute of Technology
Patrick Fulton.	Penn State University	Stephen Hickman	U.S. Geological Survey
Gareth Funning	University of California, Berkeley	Dave Hill	U.S. Geological Survey
G		George Hilley	Stanford University
lames Gaberty	Lamont-Doherty Farth Observatory	kathleen Hodgkinson	UNAVCO
Abbiiit Gangonadhyay	University of Texas at Austin	John Hole	Virginia Tech
Stenhen Gao	University of Missouri - Rolla	Austin Holland	University of Arizona
Miriam Garcia	University of Texas at El Paso	William Holt	Stony Brook University
Ed Garnero	Arizona State University	Nicolas Houlie	UC California
Hersh Gilbert	Purdue University	Stephan Husen	ETH Zurich
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J

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IRIS-USArray
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Sandra Laursen University of Colorado	
Thorne Lay University of California, Santa Cruz	

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Chris Marone	Penn State
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Glen S. Mattioli	University of Arkansas
Margaret-Ann Mayer	Dine College
Patricia McCrory	U.S. Geological Survey
Neil McGlashan	Cornell University
Jason McKenna	U.S. Army Engineer R&D
George McMechan	University of Texas At Dallas
Marcia McNutt	Monterey Bay Aquarium Research Institute
John McRaney	University of Southern California
Charles Meertens	UNAVCO
Tim Melbourne	CWU
Anne Meltzer	Lehigh University
David Mencin	UNAVCO
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Jean Miller	IAGT
Brian Mitchell	Saint Louis University
Walter Mooney	U.S. Geological Survey

13 .

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SAN ANDREAS AND CALAVERAS FAULTS FIELD TRIP

MARCH 27, 2007

Field trip leaders: David P. Schwartz, Ingrid A. Johanson, Heidi D. Stenner, Tom Fumal • US Geological Survey, Menlo Park

The San Francisco Bay region is an amazing natural laboratory for studying deformation associated with a major strike-slip plate boundary. Deformation across the Pacific-North American plate boundary in the greater San Francisco Bay region is accommodated primarily by slip on a series of right lateral strike-slip faults that includes the San Gregorio, San Andreas, Hayward, Rodgers Creek, Calaveras, Greenville, and Concord-Green Valley (Figure 1). The Bay Area has the highest density of active faults per km2 of any urban center in the United States, and the Working Group on California Earthquake Probabilities (WGCEP, 2003) has estimated a 62% likelihood of one or more $M \ge 6.7$ earthquakes on this fault system in the urban corridor between 2002 and 2031. The rate of slip across the plate boundary appears to be relatively constant at different time scales: GPS rates of 39.8 ± 1.2 mm/yr (Prescott et al., 2001) and 37.9 ± 0.6 mm/yr (d'Alessio et al., 2005); geologic slip rates on faults of 40 ± 5 mm/yr for hundreds to several thousands of years (WGCEP, 2003); and a long term (5 my) rate of 41± 1.0 mm/yr from global plate motion models (DeMets and Dixon, 1999). In contrast, regional convergence across the boundary is negligible. However, changes in strike along faults

ROAD LOG

Leave Portola Plaza hotel.

1.0 miles Merger onto CA-1 North.

- 13.4 miles Merge onto CA-156 E via Exit 414B toward Castroville/US-101/San Jose.
- 6.4 miles Merge onto US-101 N/CA-156 E toward Hollister/San Francisco.
- 8.8 miles Merge onto CA-156 E toward San Juan Bautista/Hollister.
- 2.2 miles Turn left onto Monterey St and enter San Juan Bautista.

The Mission San Juan, which you can see to the right, was founded on June 24, 1797 by Father Fermin de Lausen. It sits on the crest of a northeast-facing scarp of the San Andreas fault. The Mission is the 15th of the 21 California missions. Construction of the main church was started in 1803 and, despite damage from several earthquakes, it has been in continuous use since July 1, 1812. In addition, the Mission played key role in the conclusion of Alfred Hitchcock's 1958 film classic Vertigo that starred Jimmy Stewart and Kim Novack.

- 0.0 miles Turn left on First Street.
- 0.6 miles STOP 1. San Andreas Fault, Nyland Ranch
- 1.3 miles Backtrack through San Juan Bautista and make right onto CA-156.
- 8.3 miles Turn left onto CA-25/Bolsa Rd.
- 3.7 miles Turn right onto Shore Road.
- 1.2 miles Turn right onto Frazier Lake Road (hard packed dirt ranch road).
- 1.0 miles **STOP 2.** Calaveras Fault, Costa Ranch paleoseismic site. From here we will walk a short distance to the Calaveras fault.
- 2.2 miles Backtrack and make left onto CA-25/Bolsa Rd.
- 6.9 miles Right onto 3rd St.
- <0.1 miles Right onto West St., which becomes Virginia St.
- 0.1 miles Right on Locust St.
- <0.1 miles **STOP 3.** Calaveras Fault, Surface Creep in Hollister. Locust St. and Central Ave. From here walk southeast along the Calaveras fault to Dunne Park (approximately 500m). Buses will be waiting with lunch.
- 39.9 miles Return to Monterey.

have produced zones of transpression with associated active thrust faulting (often blind) and uplift. The most recent and prominent example of this is the 1989 Loma Prieta earthquake. This event occurred on a blind reverse-oblique slip fault beneath the Santa Cruz Mountains. which is located within a restraining bend on the San Andreas fault. Another major blind thrust occurs beneath the region's second highest peak, Mt. Diablo (3849 ft/1173 m), located in a restraining step between the Greenville and Concord-Green Valley faults (Figure 1).

In addition to locked sections of faults that fail coseismically to produce large earthquakes, the Bay Area is also home to the largest concentration of creeping faults known worldwide. These include



Figure 1. Major faults within the Pacific-North American plate boundary in the greater San Francisco Bay region. Locations of 1868 Hayward and 1989 Loma Prieta ruptures are shown in yellow; the entire length of the northern San Andreas fault (470 *km*) *ruptured in* 1906 (from WGCEP, 2003).



Figure 2. Geologic map of the San Andreas and Calaveras fault zones in the vicinity of San Juan Bautista and Hollister (Wagner et al., 2002). Field stops are shown: 1) San Andreas fault, Nyland ranch; 2) Calaveras fault, Costa ranch; 3) Calaveras fault, Hollister

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the Hayward, central and southern segments of the Calaveras, the Concord-Green Valley, the Maacama, the Sargent, as well as the southernmost part of the 1906 San Andreas rupture. The ratio of creep rate to the longer-term geologic slip rate can vary from approximately 50 percent (Hayward fault) to essentially 100 percent (central and southern Calaveras fault). An important question is: to what degree and how does creep affect the occurrence of large earthquakes on faults?

The Earthscope field trip has three stops where we will look at the San Andreas and Calaveras faults (Figure 2). These are in the vicinity of San Juan Bautista and Hollister, where the two faults accommodate about 80 percent of Pacific-North American plate boundary slip. At the latitude of San Juan Bautista the San Andreas and Calaveras faults are 11 km apart. South of Hollister the two faults converge: the Calaveras is re-named the Paicines fault and with the separation distance between the two is on approximately 3 km at the surface. The crust between the two faults in the region of convergence is complexly deforming, with seismicity distributed on active structures at depth that are not clearly recognizable at the surface (Figure 3). Other notable fault intersections/splays in the Bay Area associated with significant partitioning of slip and changes in slip rate are the Calaveras/Hayward and San Andreas/San Gregorio (Figure 1). The field trip provides the opportunity to discuss how GPS, InSAR, creep measurements, historical seismicity, and paleoseismology are being used to improve our understanding of the behavior of these major active faults.

STOP 1. San Andreas fault, Nyland Ranch (Schwartz, Johanson, Fumal)

~USE EXTREME CAUTION IN CROSSING THE ROAD!~

At this location the San Andreas fault is geomorphically expressed as a subtle, linear trough located on the west side of First Street/Old San Juan Highway. It crosses beneath an asphalt ranch road, which is cracked and right laterally offset by creep across the main trace, as are the fences on both sides. Note the narrow zone of deformation with most of the slip at the main fault trace and a small percentage of slip occurring as warping that extends only a few meters from the main trace. This narrow zone of deformation is typical of the surface expression of strike-slip ruptures, whether the offset is one meter or nine meters. Creep at this site is currently being monitored with a nail array by UC Santa Barbara.

In 1906 the fault produced surface rupture here, likely in the form of a mole-track; the amount of 1906 slip is not



Figure 3. Seismicity map and cross section showing double-difference relocated earthquakes from 1984-2000 in the San Juan Bautista (SJB)-Hollister (H) area (from Simpson et al., 2004). The vertical cross-section at top is centered on -155 (on map). Earthquakes plotted in black fall within +/-1.25 km of the vertical section. Earthquakes plotted in blue fall within +/-2.5 km of the vertical section. Red/green lines on the maps and red/green dots at the top of profile plots are faults: SA is San Andreas, C is Calaveras, and QS is Quien Sabe.

Mr Nyland reports that he has had trouble with pipe breakage at the fault at various times in the past. When asked where the 1906 earthquake fault breakage was located, he pointed at the current zone of slippage and said "I was eleven years old and living here at the time of the 1906 earthquake and remember a mound of dirt along the fault right where the road is now cracking. The mound of dirt occurred within a broad trough that was about 1m deep and was later filled to level the ground. The mound also crossed the main road to the south. I helped my father dig away the mound formed in the earthquake so that people could use the main road into San Juan Bautista again".

known, although 3.5 m of offset occurred 15 km to the north (see below). Arthur Nyland, the original owner of the ranch, was interviewed by R. Nason (1971):

Nyland ranch is on the San Juan Bautista (SJB) segment of the San Andreas fault, which exhibits complex behavior with heterogeneous slip patterns in both time and space. This section of the fault forms the transition zone between the creeping section of the San Andreas to the south and the locked Santa Cruz Mountains segment to the north. It has been recently characterized by Johanson and Burgmann (2005) using a joint inversion of InSAR and GPS data (Figure 4 and Figure 5). The SJB segment experiences a variety of transient phenomena including shallow creep events and deeper slow earthquakes (Figure 6), and moderate sized (~M5) earthquakes. Creep events are shallow slip events (<500 m), perhaps occurring when sediments reach a yield point after loading from steady creep below (Wesson, 1988). At San Juan Bautista they are generally observable on only one instrument as they involve slip on a small fault area (Gladwin, 1994). Slow earthquakes include larger and deeper fault areas than creep events and are observable on more than one instrument. Four such events have occurred (in 1992, 1996, 1998, and 2004) since strain- and creep-meter monitoring began, ranging in equivalent magnitude from 4.9 to

5.1 and with a time-span of about one week (Linde

et al., 1996, Gwyther, 2000, Bilham, 2004). The









Figure 4: Stack of nine interferograms spanning the period 1995-2001 (Johanson and Burgmann, 2005). Main figure is a close-up of the Hollister/ San Juan Bautista area; inset figure is the complete satellite frame extending from the Santa Clara Valley in the northwest to the creeping section in the southeast. Colors indicate the change in distance between the ground and the satellite in the satellite's look direction (23° off vertical). The broad deformation pattern associated with strain accumulation on the San Andreas Fault (SAF) system is visible as the frame-wide progression from blue to yellow colors. Uplift due to groundwater recharge is apparent in the Santa Clara Valley, near Hollister, and near Watsonville (white circles). Surface creep along the SAF is indicated by the sharp contrast in color from blue to green at the surface trace.

largest earthquake on this segment since a M5.5 event in 1961, was the 1998 Mw 5.1 San Juan Bautista earthquake (Uhrhammer et al., 1999). This event was immediately followed by a slow earthquake with comparable magnitude (Mw 5.0). Not only are slip rates on this segment highly time-dependent, but the secular component is heterogeneously distributed in space (Figure 5). Superimposed on a general decrease in creep rate from south to north are two low-slip/locked asperities (red dashed lines in figure 5). These low slip asperities exhibit a slip deficit that could produce an M6.3 - 6.7earthquake once a century (Johanson and Burgmann, 2005).



Figure 6: Relationship of creep events as observable from creepmeter records (top) to slow earthquakes as observable in strainmeter records (bottom, from http://www.gtsmtechnologies.com/index_files/nehrp. htm). Creep events, being shallow phenomena, affect only the near-

fault creepmeter. However, slow earthquakes, which involve slip on a larger fault area, are discernable on strainmeters that are located off-fault.

The Mill Canyon-Arano Flat Paleoseismic Site

Fifteen kilometers to the north of STOP 1, USGS geologists under the guidance of Tom Fumal have developed two paleoseismic sites about 0.6 km apart on the San Andreas fault at Mill Canyon and Arano flat (Figure 5). These sites are on the Kelly-Thompson ranch, which has been in the same family since 1851 and has been preserved in a near pristine condition. The fault is geomorphically very well expressed along this reach. The combination of rapid deposition and abundant datable organic material has resulted in the first high-resolution chronology of large earthquakes on the Santa Cruz Mountains segment of the fault.

At Mill Canyon the most recent ground rupturing event, the 1906 San Francisco earthquake, is well-expressed as a series of in-filled fissures and small scarps. Evidence was found for three additional ground-rupturing earthquakes since about A.D. 1500 (Figure 7). Of particular interest at this site is the dating of the penultimate earthquake and its relation to the Bay Area earthquake of 1838. The 1838 event is generally considered to have occurred on the San Andreas fault. While this is a reasonable interpretation from the limited historical information there is uncertainty regarding its location and magnitude; and, at present, no direct geologic evidence of this event has been observed. Bakun (1999), based on estimated intensities from damage reports, calculated the magnitude as M6.8 \pm 0.4 and suggested the Peninsula San Andreas as the most likely source. Several investigators (Tuttle and Sykes, 1992; Toppozada and Borchardt, 1998) have speculated that the 1838 earthquake ruptured both the Peninsula and Santa Cruz Mountains segments of the



Figure 7. Photomosaic of the northwest wall of trench 6 at Mill Canyon (Fumal, in preparation). Faulting during the 1906 earthquake occurred along a single trace in this exposure producing a small scarp and accompanying small fissure. An earthquake about A.D. 1650-1730 occurred when the organic soil overlying gravelly unit 8 was at the ground surface, producing a small scarp and accompanying colluvial wedge consisting of fragment of unit 8 in an organic-rich matrix. This wedge is overlain by sand and gravel of unit 5, which was subsequently faulted about A.D. 1700-1770, producing a deep, gravel-filled fissure. Evidence for a fourth earthquake is not visible in this exposure but appears as a fissure filled with unit 8 gravel on the southeast wall of trench 6.

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San Andreas fault, producing an earthquake of M 7.3-7.5 with 90-130 km of surface rupture. Radiocarbon ages of detrital charcoal at Mill Canyon suggest an age of the penultimate earthquake of A.D. 1700-1776, however. Additional support for this age comes from a 1.5-meter-deep fissure that formed during this earthquake. It was sampled for Erodium, a non-native pollen that first appeared in the San Francisco Bay area about 1770 and was ubiquitous in the region by 1800. Erodium pollen would have been present at this site in 1838 and should be found in the fill of any fissure formed during an earthquake of that vintage. The presence or absence of this pollen in critical deposits is presently playing an important role in dating Bay Area paleoearthquakes that occurred between the early 1700s and early-middle 1800s, a time period in which radiocarbon gives multiple ages. Along with observations 15 km further north at Grizzly Flat (Schwartz and others, 1998), the Mill Canyon results indicate that if the 1838 event was on the San Andreas, it did not extend into Santa Cruz mountains.



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Figure 8. Photomosaic plan view of the floor of an excavation exposing a small channel offset across the San Andreas fault (Fumal, in preparation). The 1906 fault traces are in red. Three piercing points were measured: 1) the northwest margin of the channel cut into older fluvial sediments (blue), 2) the northwest edge of the upper layer of sand and gravel in the channel fill (yellow), and the southeast edge of a lower cobble gravel in the thalweg of the channel (green). Offsets range from 3.2 to 3.5 The upper sand and gravel layer contained abundant artifacts including bottles made only between 1887 and 1890.

At Arano Flat, faulting is expressed as a 1 to 2-m-wide zone that deforms alluvial fan deposits overlying well-bedded overbank deposits The trenches at this location contained evidence for at least nine surface-faulting earthquakes since about A.D. 1000. Evidence for ground-rupturing events includes in-filled fissures, folding with growth strata, and multiple upward terminations of fault traces. Earthquake ages were constrained using an OxCal chronological model incorporating AMS (accelerator mass spectrometer) radiocarbon ages of 113 samples of detrital charcoal from 19 layers and stratigraphic ordering. The mean recurrence interval is about 105 years, while individual intervals range from about 10-310 years. One of the key questions that affects the understanding of fault behavior, hazard analysis, ground motion estimates, and insurance rates is whether the short recurrence times observed here represent only large events on the Santa Cruz Mountains segment of the fault (including multiple segment ruptures such as 1906) or whether the site is also recording



Figure 9. Topographic map of the Costa Ranch site. Trenches are numbered 1-6, arrows point at main fault where a down-to-the-east scarp is formed. Contour interval is 15 cm (Stenner, in preparation).

overlapping ruptures from the San Juan Bautista segment.

Arano Flat also provides one of the few observations of 1906 slip along the southern section of the 1906 rupture. A small tributary channel of Arano Creek flowed across the alluvial flat during the 19th century, crossed the fault, and rejoined the main stream. The upper part of the fill in this channel contains abundant artifacts, mostly fragments of glass and ceramics associated with a small house that was located about 700 m southeast of the trench site during the turn of the century. These artifacts include intact bottles made only from 1887-1890. Progressive excavation of this channel exposed it across the fault (Figure 8). The channel is offset 3.2- 3.5 m. Given that that there is no evidence either here or at Mill Canyon of the 1890 M 6.3 earthquake (Bakun, 1999), which produced surface rupture on the San Andreas fault southeast of Pajaro Gap, this entire amount of slip likely occurred during the 1906 earthquake. This slip value is high compared to the geodetic estimate of 2.3-3.1 m for slip at depth (Thatcher et al., 1997) or the geologic estimate of 1.7-1.8 m of surface slip at Wright's tunnel (Prentice and Ponti, 1997), about 33 km northwest of Arano Flat.



STOP 2. Calaveras Fault, Costa Ranch Paleoseismic Site (Stenner)

Figure 10. View to the north of Trench 5 across Calaveras fault sag pond at Costa Ranch. Red dashed lines are surface traces of faults exposed in the trench. Photo by H. Stenner.

The Calaveras fault splays from the San Andreas (or joins it, depending on your point of view) south of Hollister and extends 125 km north to Danville in the San Ramon Valley. It is one of the most active and complex faults in the Bay Area. The northern 40 km (referred to as the Northern Calaveras fault) has little to no creep, is essentially devoid of microearthquakes, and is considered capable of producing M 6.8-6.9 earthquakes (WGCEP, 2003). In contrast, the central/ southern Calaveras fault is characterized by abundant microearthquake activity that likely reflects the high rate of creep. The average creep rate from 1968 to 1999 was 16.3 mm/yr (Galehouse and Lienkaemper, 2003); the average geologic slip rate is 14 ± 5 mm/yr for the past 4000 years (Kelson et al., 2001). Within the uncertainty of the geologic data, the long-term slip rate on the central Calaveras is consistent with the short-term slip rate derived from aseismic slip data and geodetic modeling. Historically, the central Calaveras fault has been the source of moderate magnitude earthquakes (1949 Gilroy, M 5.2; Coyote lake, M 5.9; 1984 Morgan Hill, M 6.2; 1988 Alum Rock, M 5.1). A similar sequence of comparably sized earthquakes occurred between 1897 and 1911. Based on the historical seismic record and the agreement between geodetic/creep rates and geologic slip rates, the WGCEP (2003) concluded that there is little or no strain accumulation that would result in a large earthquake along this section of the fault. However, evidence from trenches at San Ysidro Creek on the central Calaveras can be interpreted as support for ≥ 2 m-displacement surface rupturing earthquakes, with three events in the past four thousand years (Kelson et al., 2001).

At the Costa Ranch site we will walk a 300m-long section of the Calaveras fault that exhibits classical strike-slip geomorphology characterized by fault scarps and sag ponds (Figure 9). A fundamental question is whether these surface features can be formed by creep alone, or also require coseismic surface rupture during large earthquakes. From 1971-79, a creepmeter at this location recorded 15 mm/yr of creep, and another creepmeter at Shore Road, less than



Figure 11. Log of south wall of Trench 6 at Costa Ranch on the Calaveras fault showing structure and stratigraphy of the northern sag pond. The sag is bounded on both eastern and western sides by obliquely slipping faults that allow the section between to subside. The main fault zone, which accommodates both normal and lateral slip, is located immediately east of the pond's middle. The position of this trace suggests that a majority of slip occurs through the extensional sag pond area on at least one major Reidel shear, possibly acting as a linking structure as the overall fault geometry bends to the right. Minor faulting occurs throughout the sag pond, about 6 meters wide, but the main shear zone is ~1 m across. The faults exposed in the trench either reach the surface or, as within the fairly massive sag pond clay, are difficult to trace upward (Stenner, in preparation).

2 km to the north, recorded 6.5-12 mm/yr, depending on the averaging technique (Schulz and others, 1979; Schulz and others, 1982). The fault trace makes a 25-30 m right step at the southern end of the site, and at the north end of the site it bends to the east, producing sag ponds in the resulting extensional zones at both locations (Figure 9). In total six trenches were excavated at the Costa Ranch (unfortunately none is presently open). Trenches 1 and 2 crossed the large step-over; Trenches 3-6 crossed the fault scarp and sag pond to the north (Figures 10 and 11).

Trenches 1 and 2 exposed 1 to 2.5 m of overbank sediments (silt, sand, and clay) on which a moderately to very heavily bioturbated organic horizon has developed. No evidence for distinct colluvial wedge packages or fissures (both of which are indicators of instantaneous surface rupture that would be associated with large earthquakes), or consistent upward fault terminations were observed. Most of the faults were traced to the surface because of continuous creep, and some strands were obscured in the bioturbated soil horizon. Trenches 3 and 4 exposed stratigraphy similar to that in Trenches 1 and 2: fluvially deposited silt, sand, and clay overlain by an organic horizon. Three radiocarbon dates within the upper fluvial sediments yielded a range of 1800–2700 years BP. In Trenches 3 and 4 some faults extend to the surface as obviously creeping traces, and others are likely creeping but have not experienced sufficient creep to demarcate their location through the young, bioturbated upper soil horizon. All deformation observed in Trenches 3 and 4 can be explained solely by creep processes, but coseismic rupture is not completely excluded.

Trench 5 and Trench 6 were excavated across the northern sag pond (Figure 9). A photograph of the north wall of Trench 5 and a log of the south wall of Trench 6 are shown on Figures 10 and 11, respectively. Both trenches exposed the youngest stratigraphy at Costa Ranch, with more than a meter of accumulated sag pond sediment carbon-dated at younger than 1000 yrBP. Two dated shells suggest an even younger deposit of 300 yrBP (Trench 5). The sag pond strata, observed in both Trenches 5 and 6, are progressively warped down into the sag and thicken towards the center. The uppermost unit, likely historical, also thickens into the center of the sag, and is faulted (vertical component ~1 cm) by at least two creeping traces. The base of the sag pond was not exposed in either trench. The total vertical component of slip is therefore more than 2 meters. Using an estimated range in age for the oldest sag pond sediment of 1000-2000 yBP, a minimum vertical slip rate would be 1-2 mm/yr. There is no geologic basis for independently estimating the horizontal slip rate at this location.

In summary, no evidence for coseismic surface rupture of significant size has been found
on the 11 trench walls exposing the southern

Calaveras fault in late Holocene sediment at Costa Ranch. Fissures fills or consistent upward terminations were not observed in any of the trenches. Creep and micro-to-moderate magnitude seismicity may accommodate all of the fault slip, resulting in a low probability of future large earthquakes rupturing the central and southern parts of the fault. The lack of evidence cannot preclude the possibility, however, that the fault has ruptured in a large earthquake. With the high creep rate for this section of the fault evidence for large, coseismic surface rupture may have been overprinted and modified by creep. Large coseismic rupture may also have occurred at depth and did not reach the surface, or a rupture may have been severely attenuated toward the surface because of creep.

STOP 3. Calaveras Fault, Dunne Park Area, Hollister (*Stenner, Schwartz*)

In contrast to Stop 2, where we observed the long-term geologic effects of creep in unconsolidated deposits, we can see the effect of creep on manmade structures including roads, curbs, concrete retaining walls, and houses. Surface creep and its effects can be observed in urbanized areas along other Bay Area faults including the Hayward (especially in Fremont, Hayward, and Berkeley), Concord-Green Valley (in Concord),



and Maacama (great creep in Willits). However, the Hollister creep features are the best developed because of the high creep rate of 12-14 mm/yr at this location (Lisowski et al., 1991). Hollister experienced significant damage from the 1906 and 1989 earthquakes, as well as minor damage from the 1979 Coyote Lake (M 5.6) and 1984 Morgan Hill (M6.2) earthquakes on the Calaveras fault. Along with Parkfield, Hollister bills itself as the "earthquake capital".

We'll walk from the intersection of Locust St. and Central Ave. (Figure 11, Stop 3A) six blocks southeast to Dunne Park (Stop 3B). The west-facing break in slope extending across the grassy section of the park is the scarp of the Calaveras fault. Weather permitting we'll eat lunch on the fault trace, and maybe experience a small event or two for dessert, before returning to Monterey.

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